



Technische
Universität
Braunschweig

Institut für Geoökologie
Bodenkunde und Bodenphysik

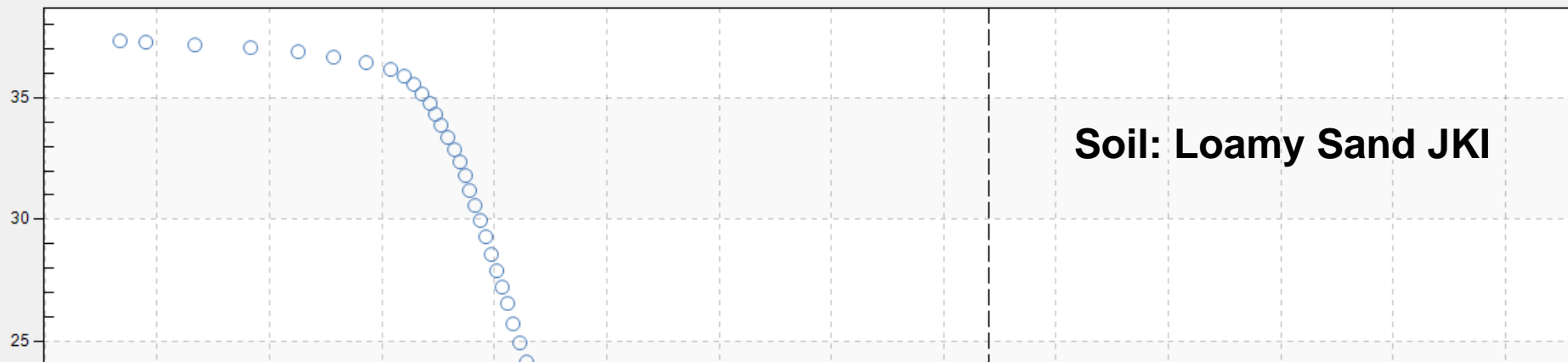


Parameterization of models for soil hydraulic properties: Challenges and recent advances

Wolfgang Durner

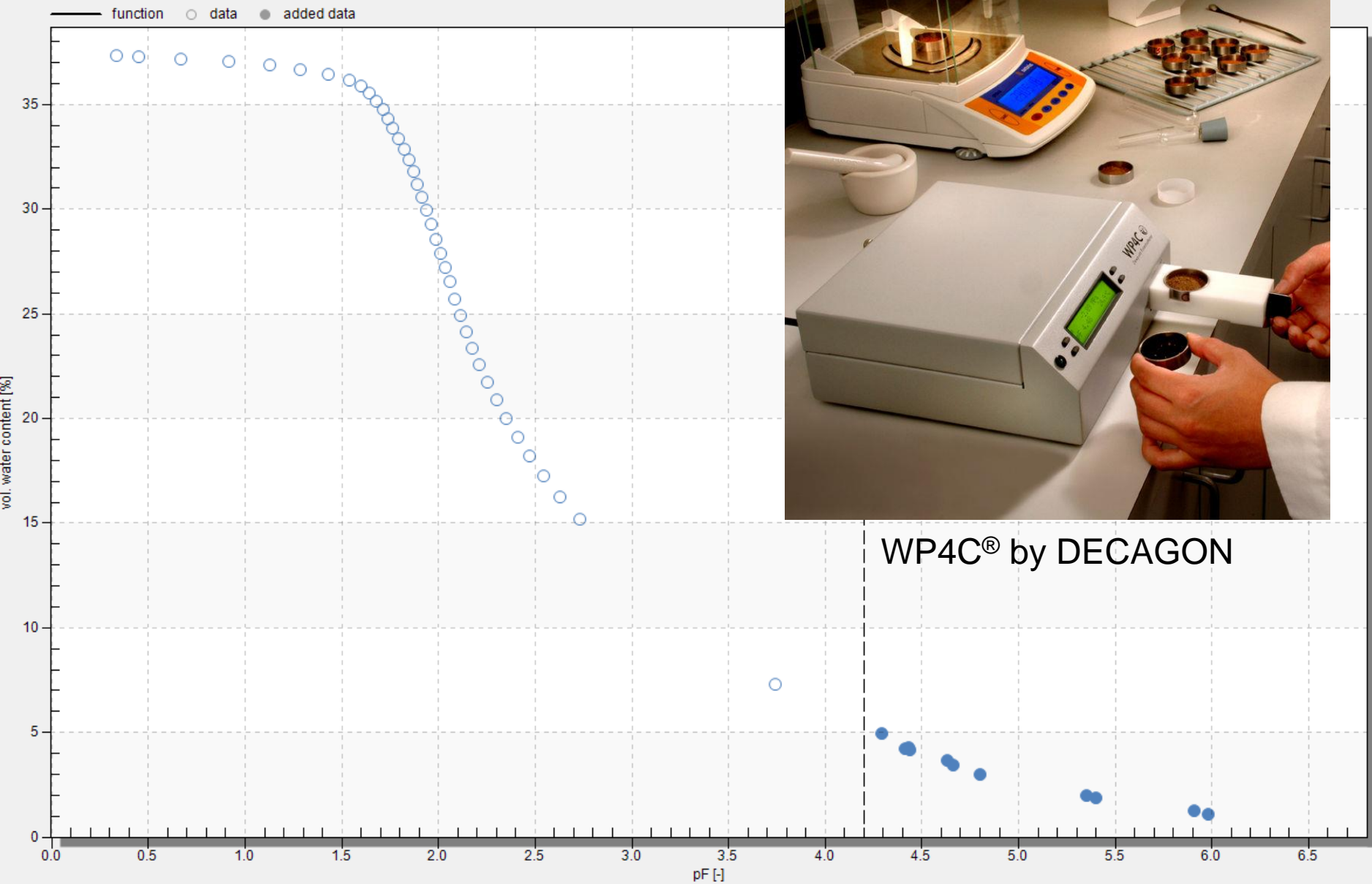
Sascha Iden, Andre Peters, Efstathios Diamantopoulos,
Kai Germer, Tobias Weber, Benedikt Scharnagl

— function ○ data ● added data



HYPROP by UMS





File Extra Help



Information Measurements Evaluation Fitting Export

Soil hydraulic model selection

	original	PDI	bimodal	bimodal PDI
Brooks-Corey	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>
Fredlund-Xing	<input type="radio"/>	<input checked="" type="radio"/>	<input type="radio"/>	<input type="radio"/>
Kosugi	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>
van Genuchten m=1-1/n	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>
van Genuchten mvar	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>

Model: FX-PDI Model Code: 6111 ?
 Source: not yet published
 Description: PDI-variant of the Fredlund-Xing model

Retention function parameters

Parameter	Value	2.5%	97.5%	Unit
<input checked="" type="checkbox"/> alpha	0.0291	0.0288	0.0294	1/cm
<input checked="" type="checkbox"/> n	15.000*	13.139	15*	-
<input checked="" type="checkbox"/> th_r	0.041	0.032	0.050	cm ³ /cm ³
<input checked="" type="checkbox"/> th_s	0.343	0.339	0.347	cm ³ /cm ³
<input type="checkbox"/> pF_dry	6.80			-
<input checked="" type="checkbox"/> m	0.506	0.418	0.595	-

Conductivity function parameters

Parameter	Value	2.5%	97.5%	Unit
<input checked="" type="checkbox"/> Ks	840.4	478.7	1475.2	cm/d
<input checked="" type="checkbox"/> tau	2.228	1.701	2.756	-
<input checked="" type="checkbox"/> omega	1.35E-5	6.59E-6	2.78E-5	-
<input type="checkbox"/> a	-1.500			-

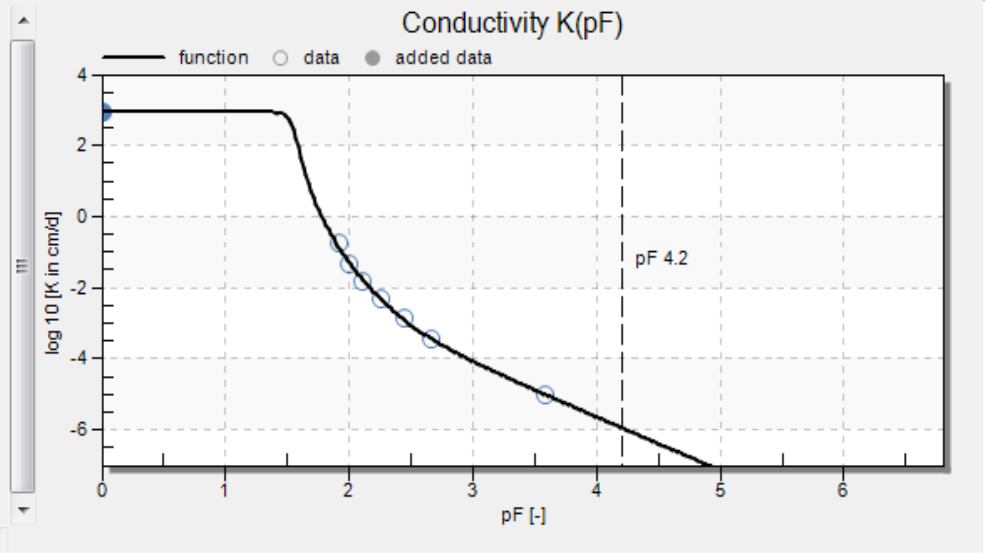
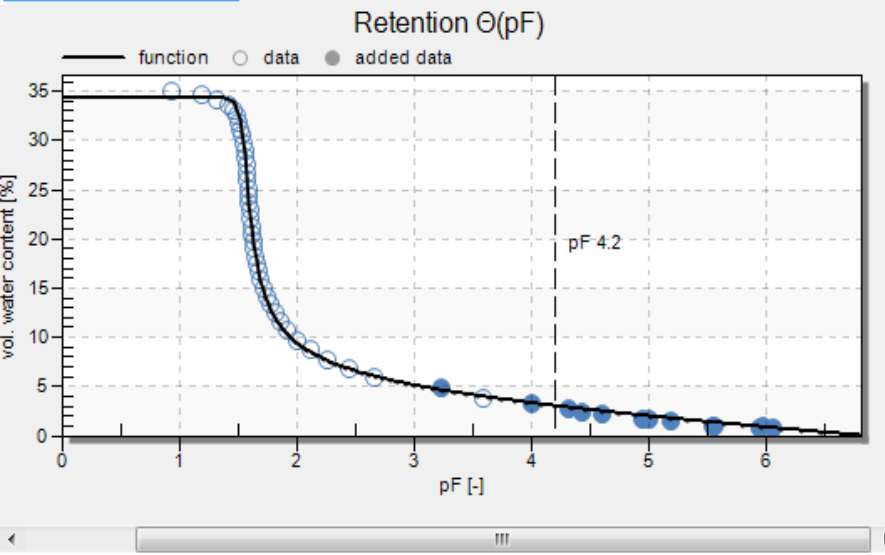
Statistical analysis

Parameter Correlation Matrix
AICc -666

Field capacity

WC @ 6 kPa 12.7 %
 WWC @ 6 kPa 12.7 %
Optimization Parameters
 PAW 0...1500 kPa 5.7 %
 PAW 33...1500 kPa 3.5 %

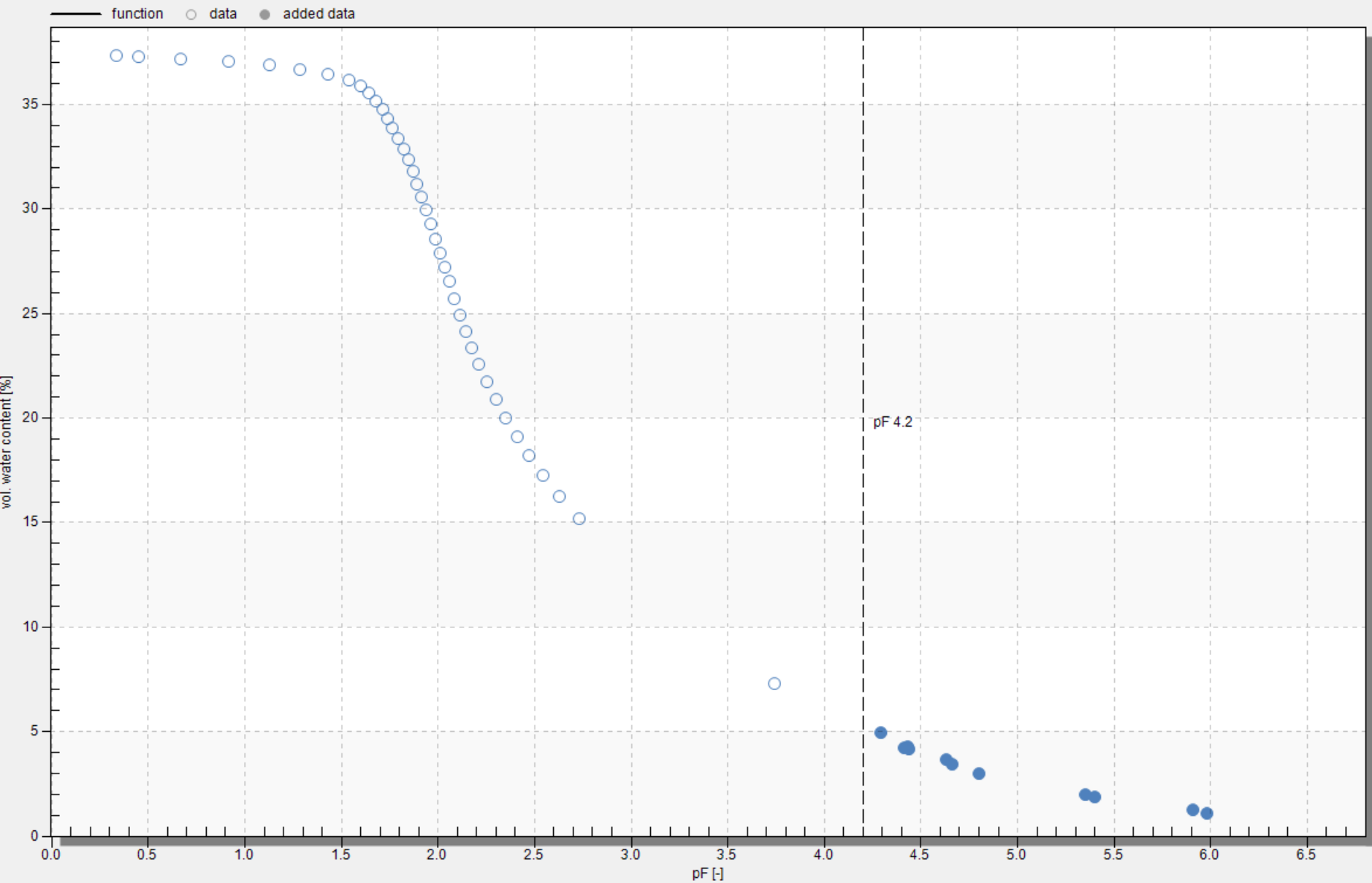
Retention Conductivity K(theta) Show data as small points

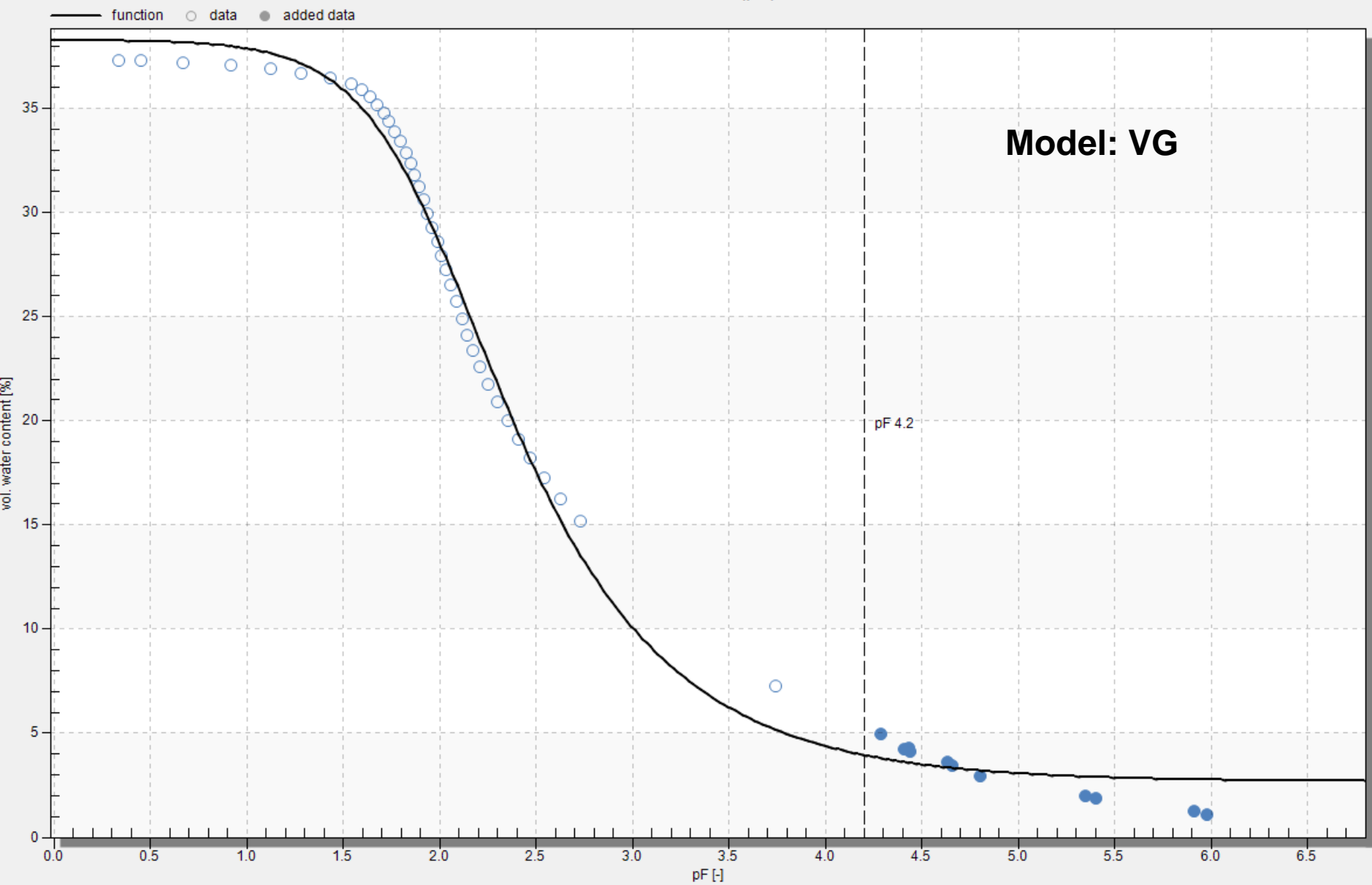


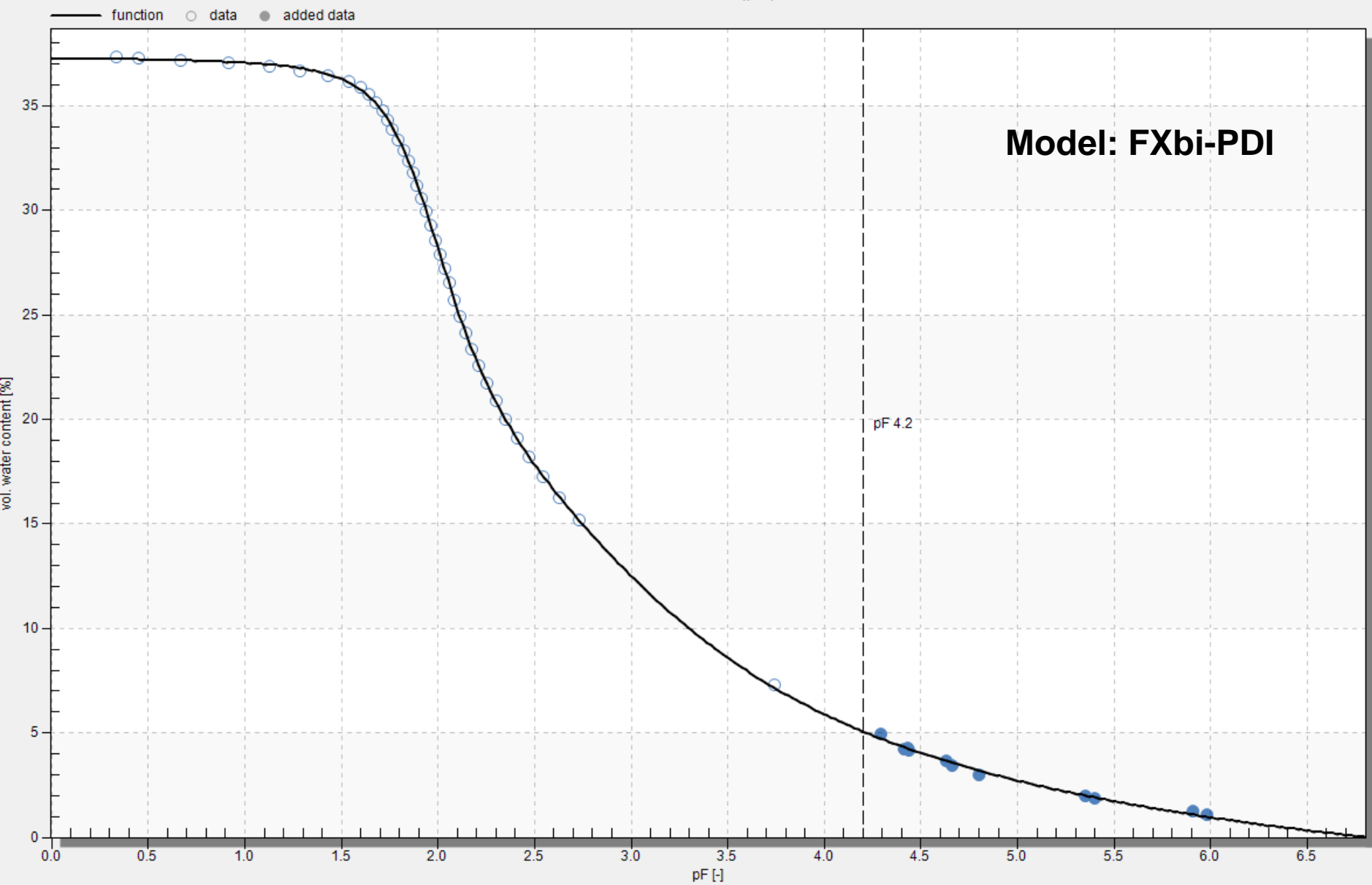
PowerUser

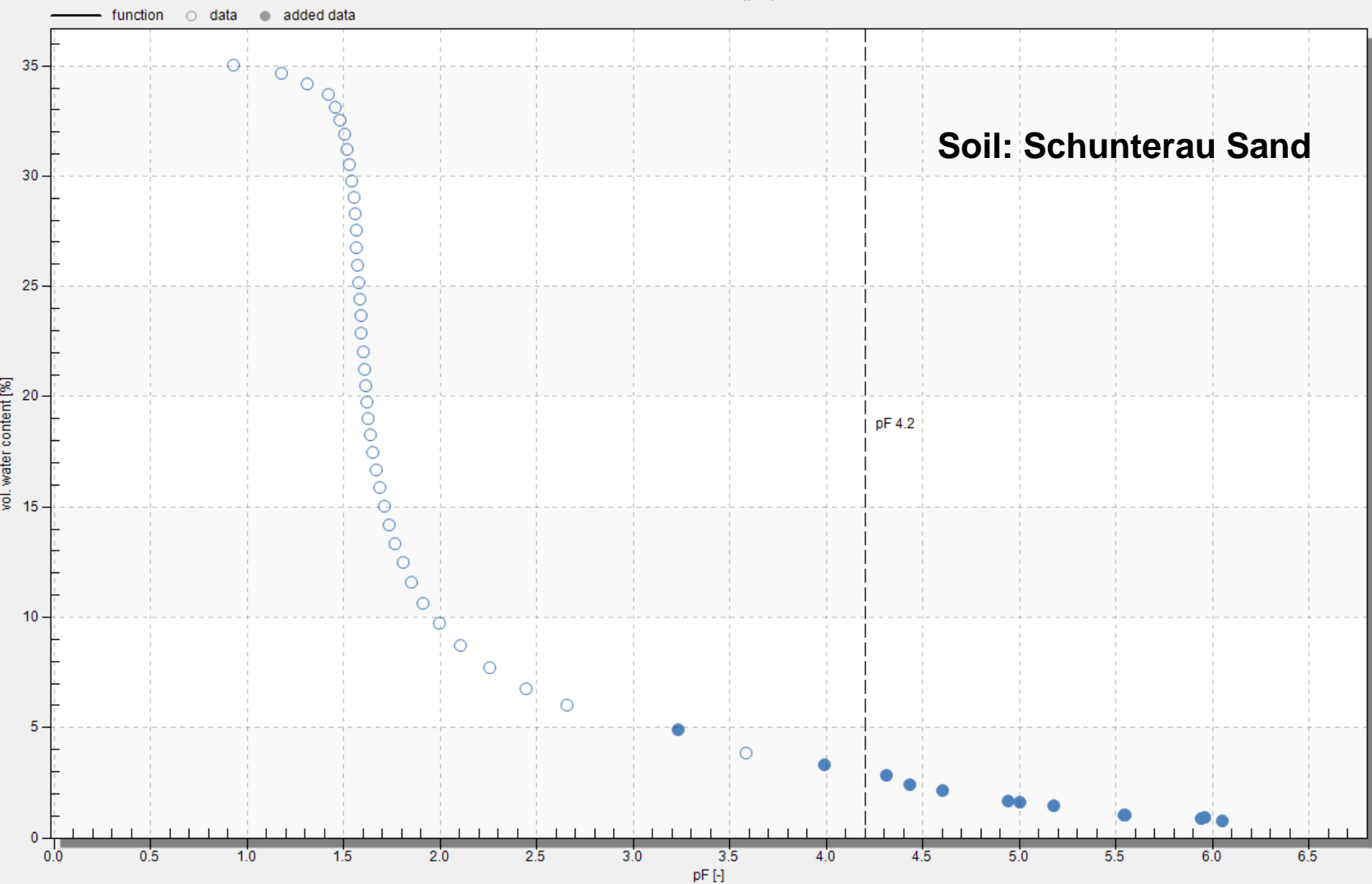


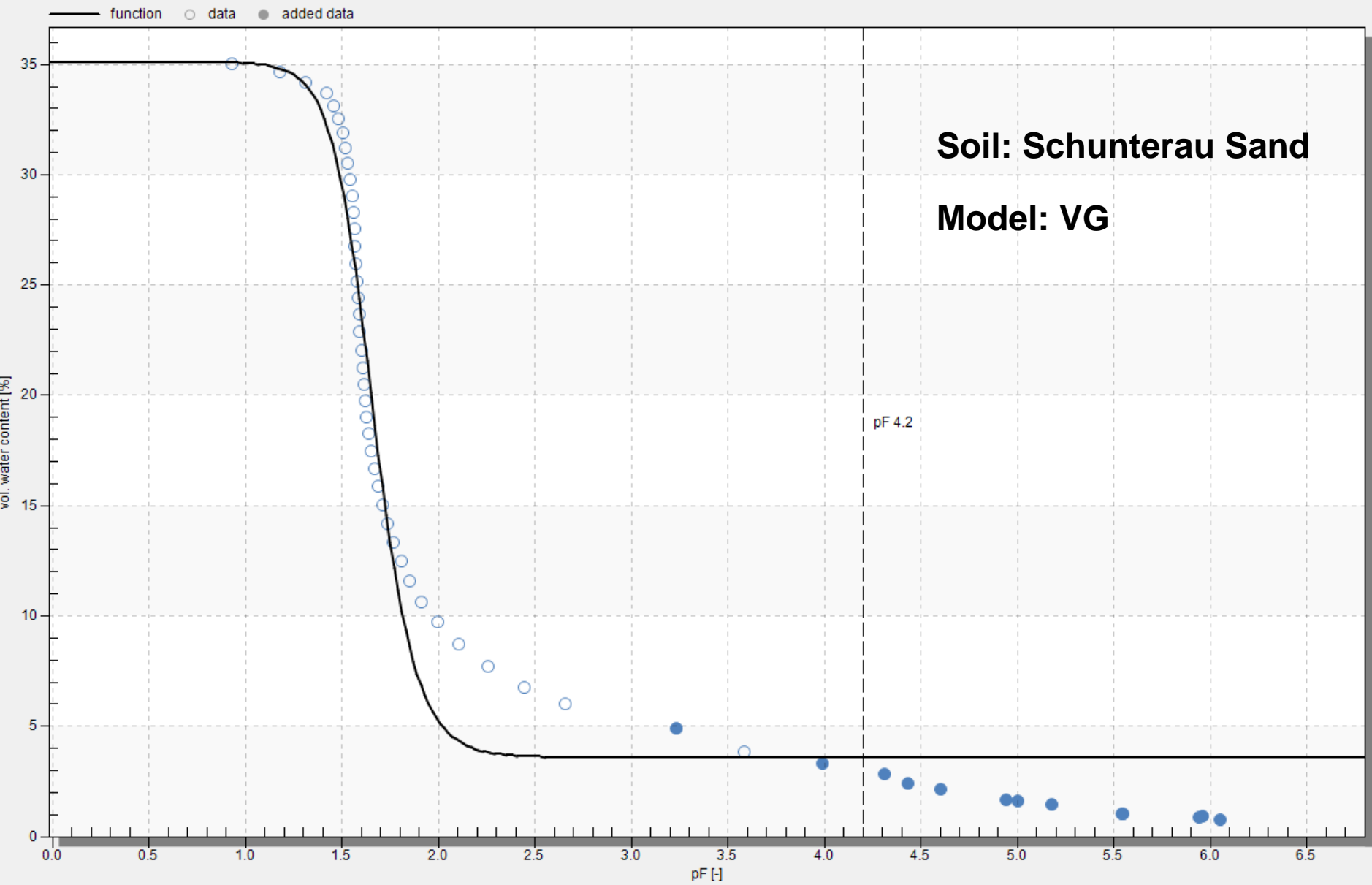
Retention $\Theta(pF)$

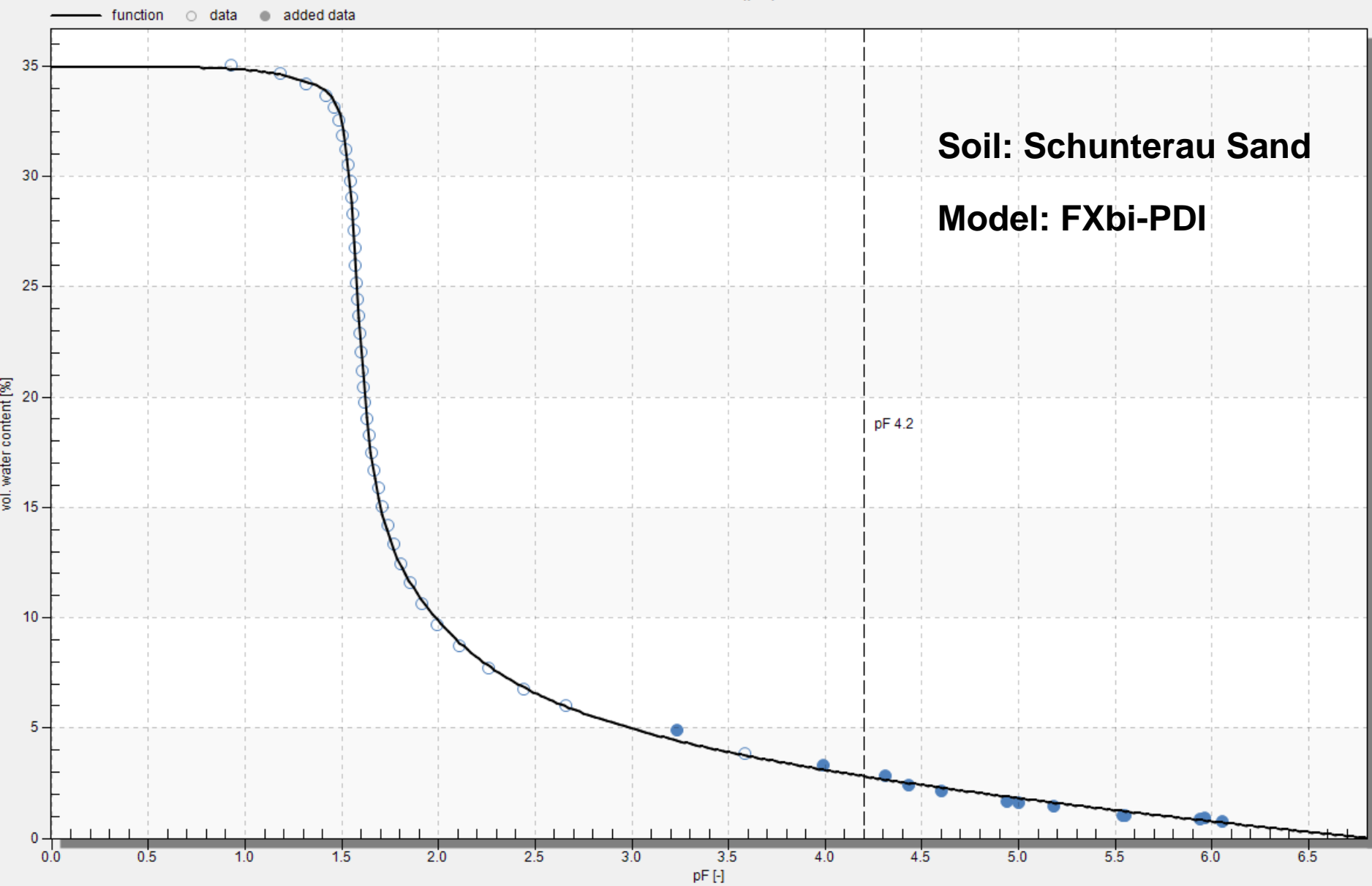






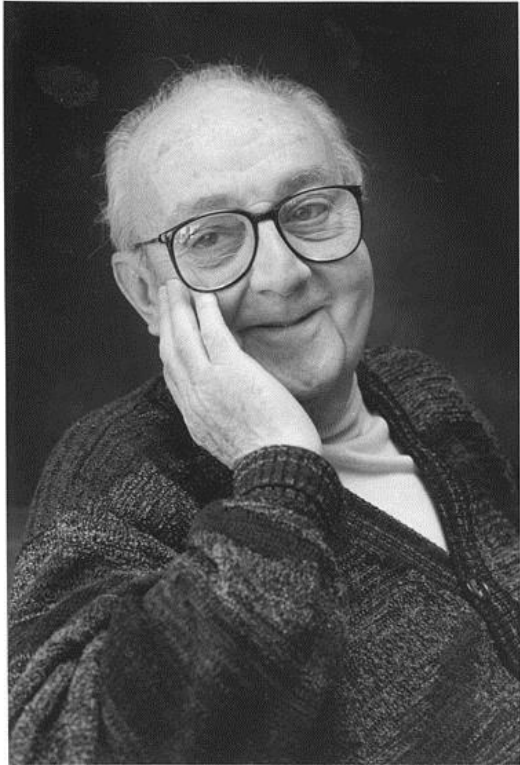






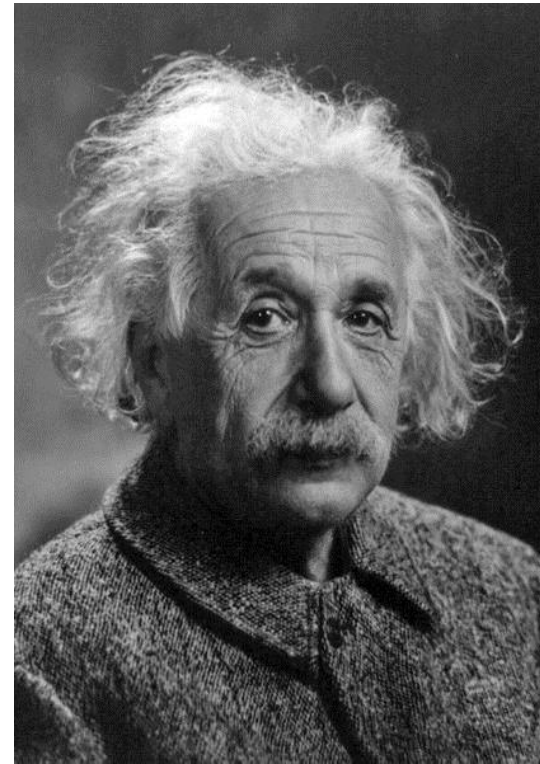
Parameterization

**„All models are wrong ...
... but some are useful“**



G.E.P. Box 1919 –2013

**„Models should be as simple as possible ...
... but not simpler“**



A. Einstein 1879 –1955



How good (accurate) should a model fit be?

... depends on the intended use of the model

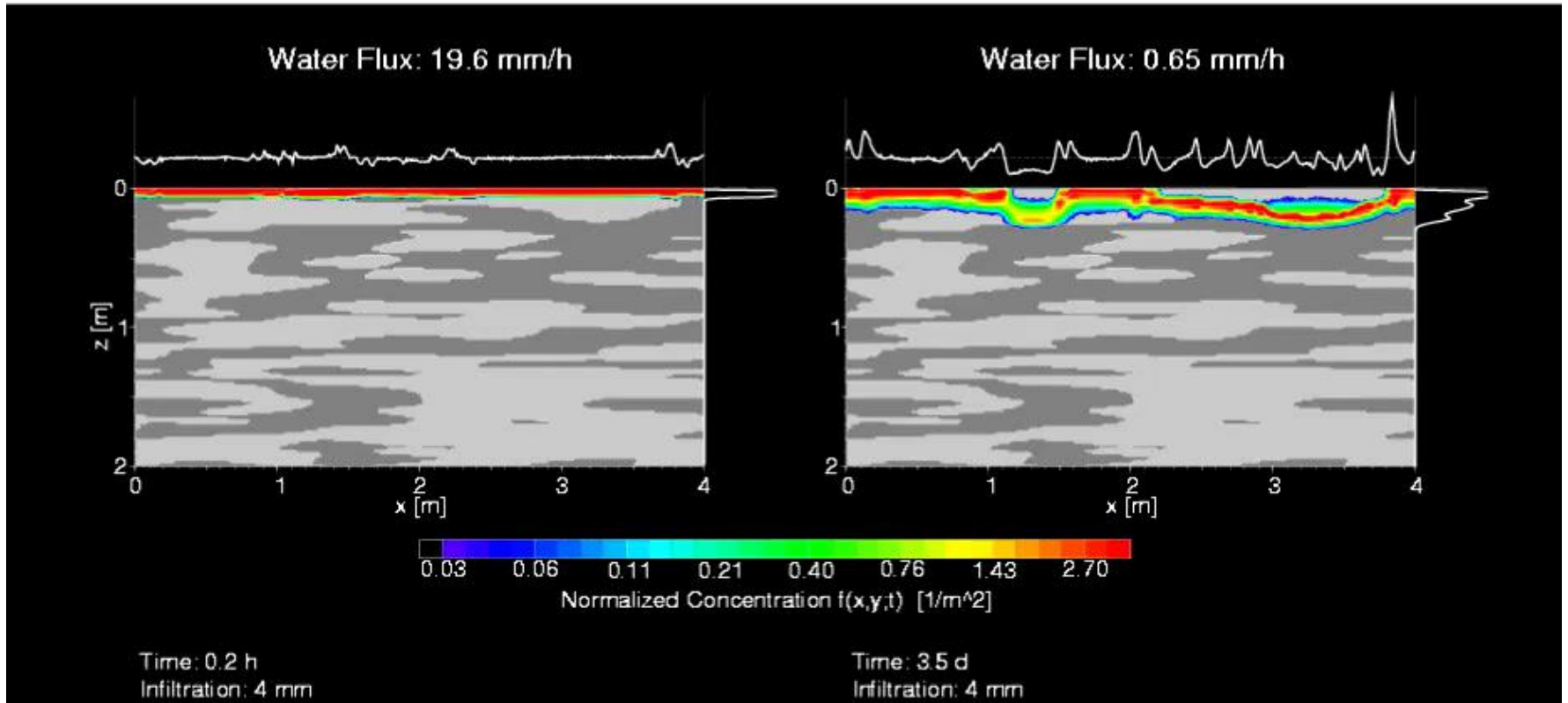
- Field capacity
- Plant available water content
- Mapping of soil hydraulic properties
- RETC to derive secondary properties, e.g., predicting hydraulic conductivity
- Modeling soil water dynamics



The working horse

**Richards Equation is today's
standard model for simulating soil
water dynamics**

(Infiltration, redistribution, evaporation,
transpiration, drainage)



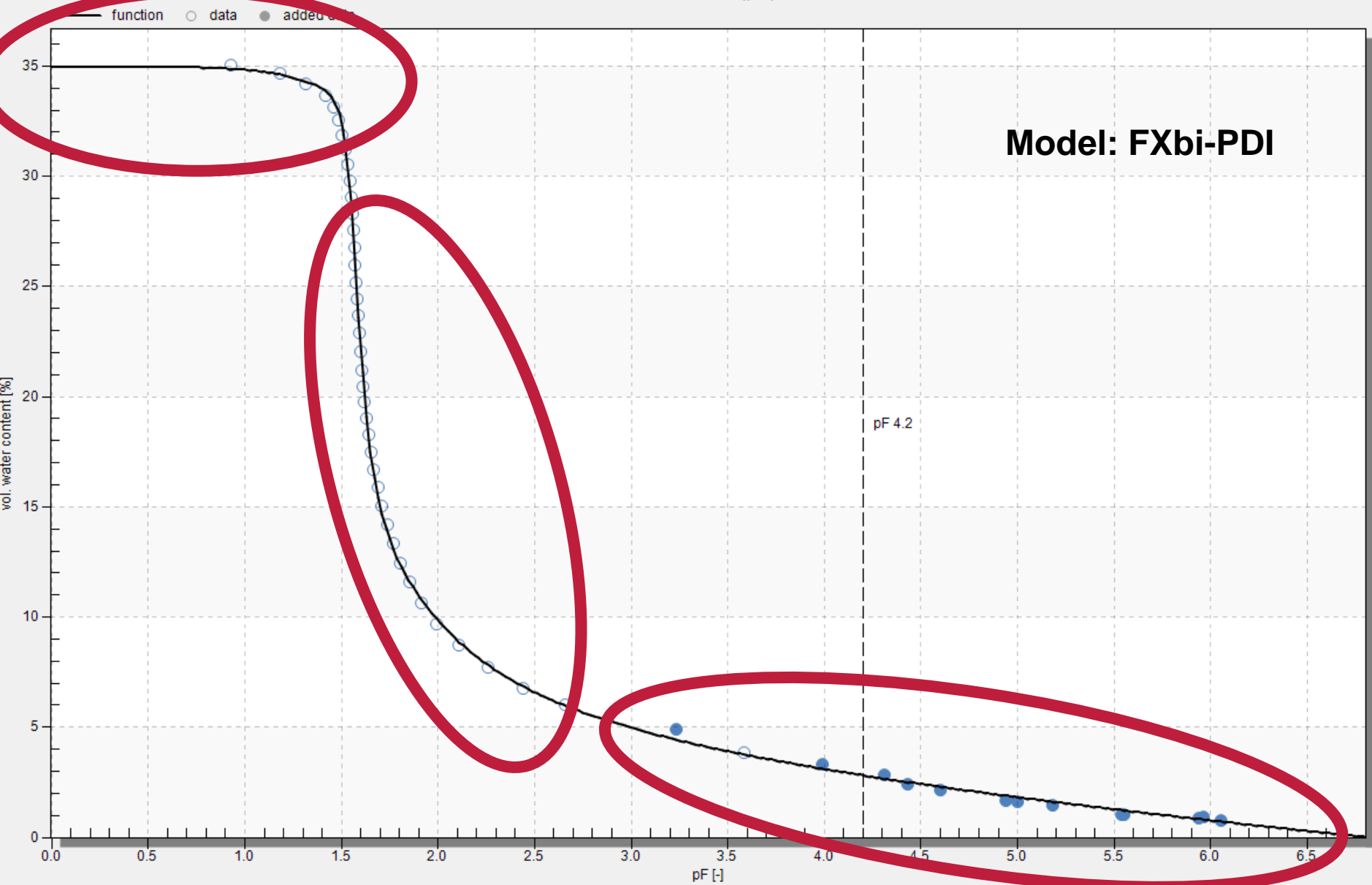
Courtesy Kurt Roth

- Extent of evaporation ?
- Water flow to roots ?



- Erosion ?
- Solute transport ?

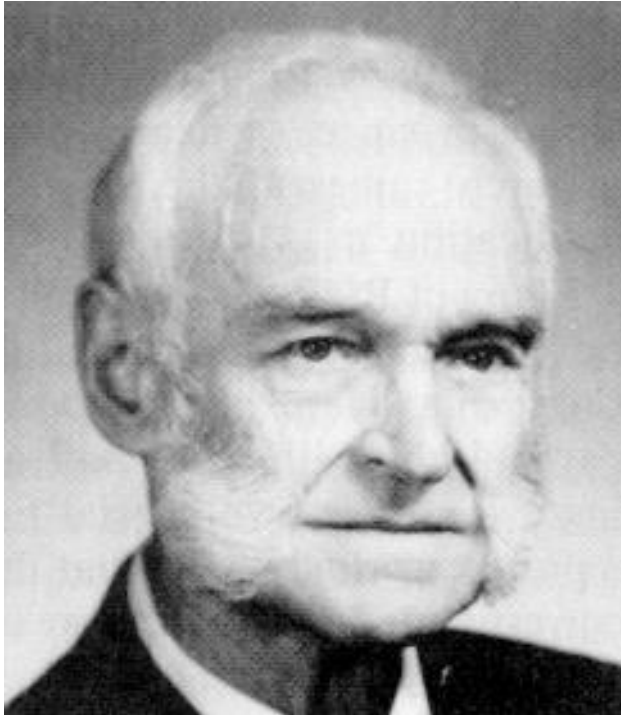






Part 2: A look on hydraulic properties

Lorenzo A. Richards (1931)



...combined **Continuum equation** with **Darcy-Buckingham equation** (here: 1D, vertical).

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[K(h) \left(\frac{\partial h}{\partial z} + 1 \right) \right] + s(z, t)$$

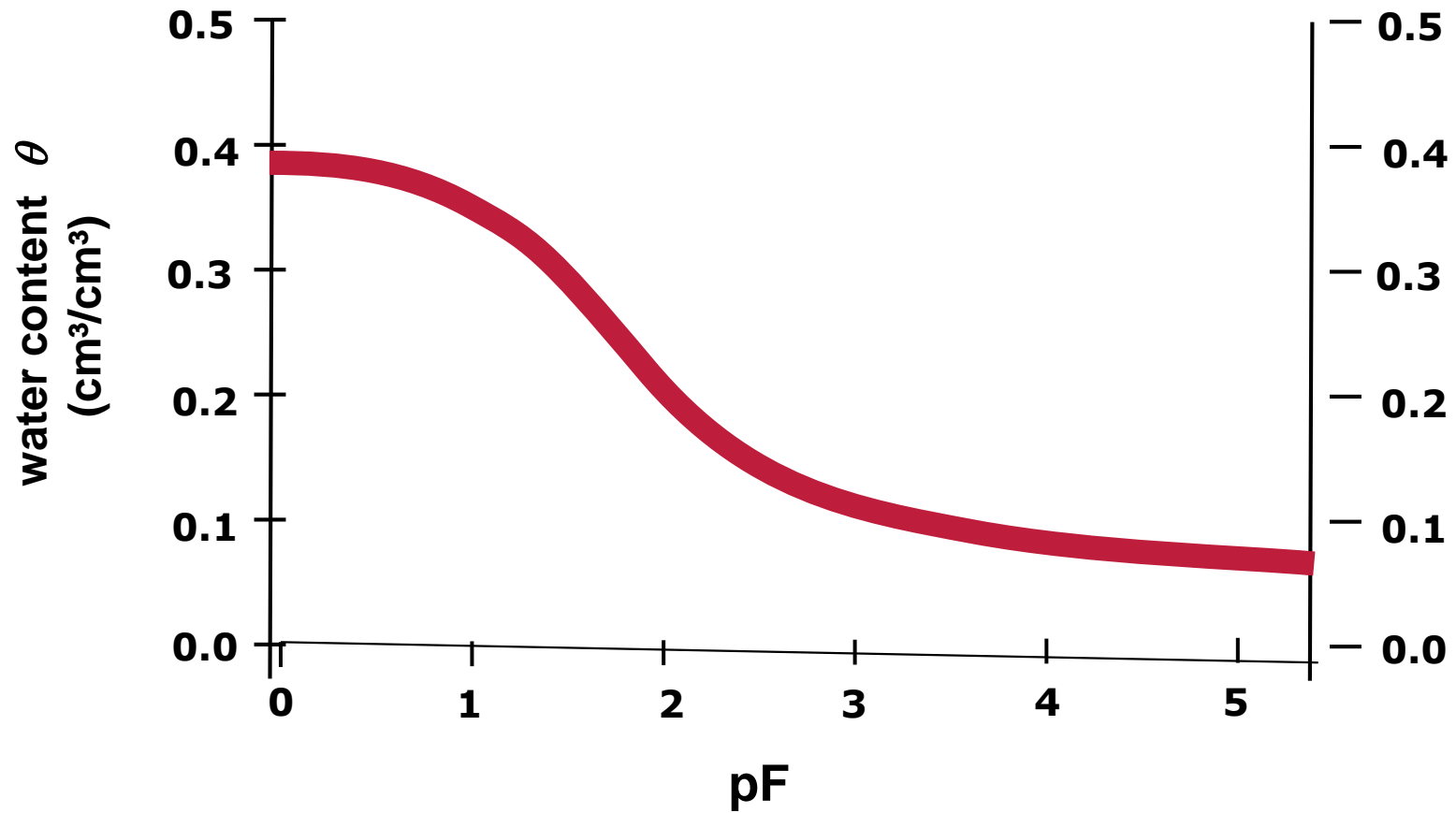
Constitutive relationships

1) Conductivity curve, $K(h)$

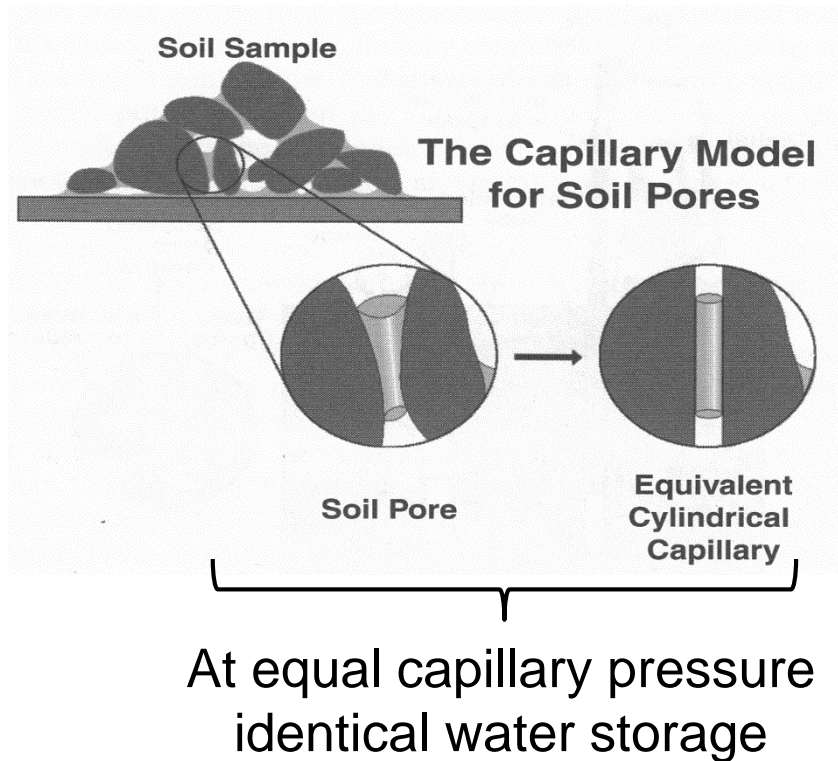
2) Retention curve, $\theta(h)$
("most important curve"
in soil physics)

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[K(h) \left(\frac{\partial h}{\partial z} + 1 \right) \right]$$

Van Genuchten retention curve



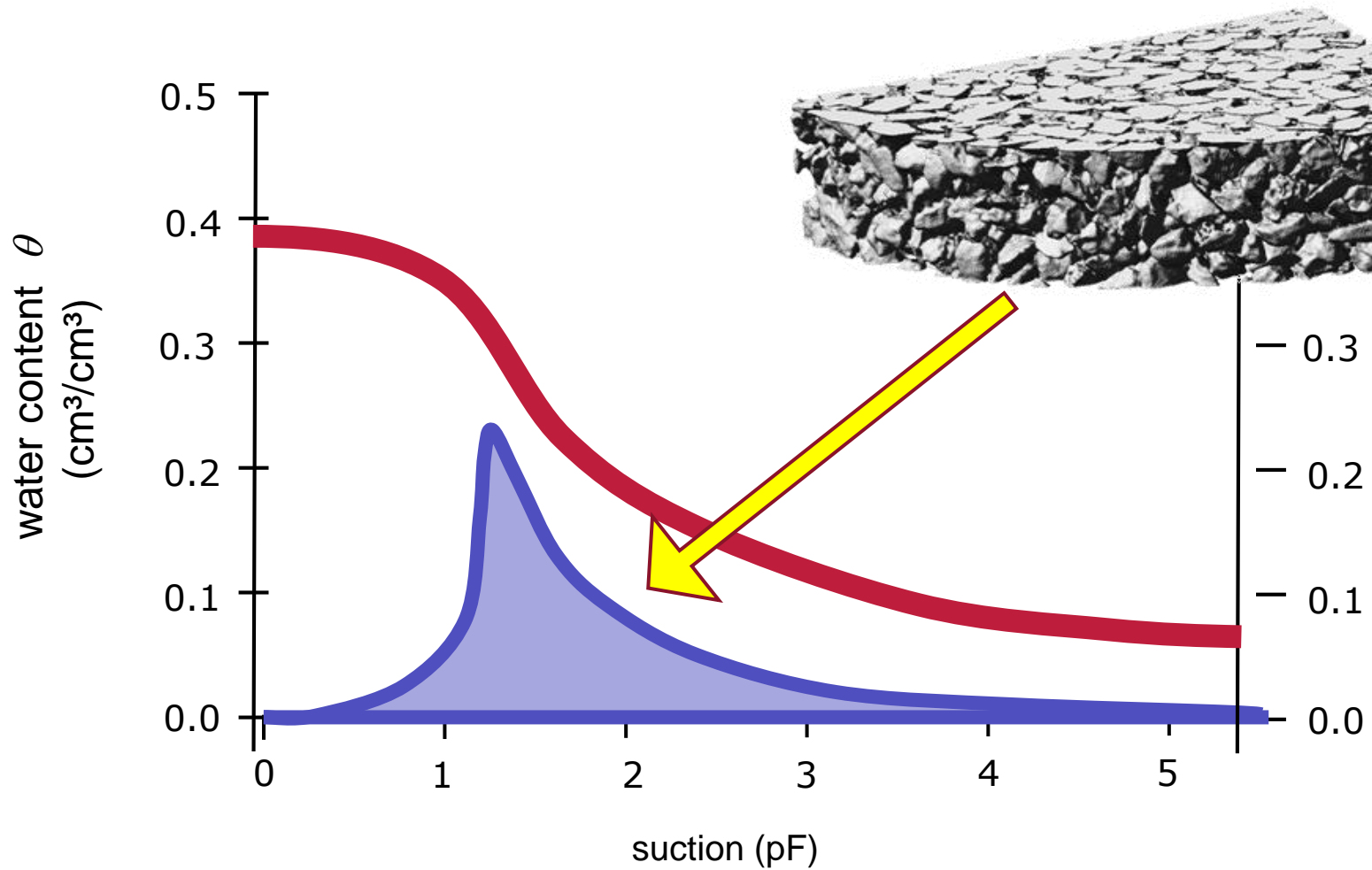
Concept of equivalent pore sizes



Concept of cylindrical equivalent pores leads to a representation of the true pore space as a „capillary bundle“.

Allows to derive pore-size distribution, water holding capacity, and hydraulic conductivity

Equivalent pore-size distribution

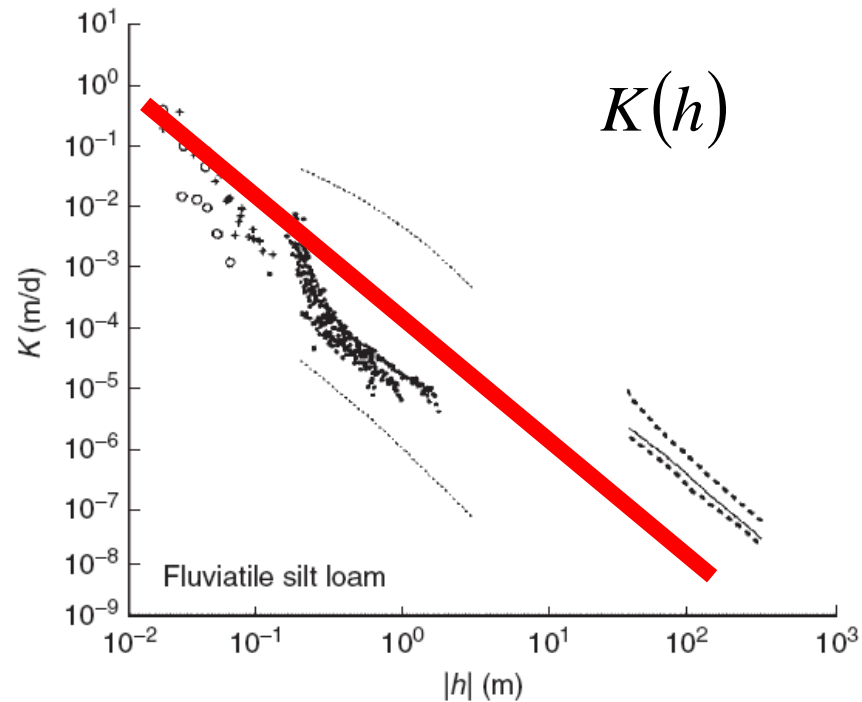


Constitutive relationships

1) Conductivity curve

2) Retention curve (most important curve in soil physics ?)

„hydraulic properties“: Hydraulic conductivity function



○ Crust method

+ Drip infiltrometer method

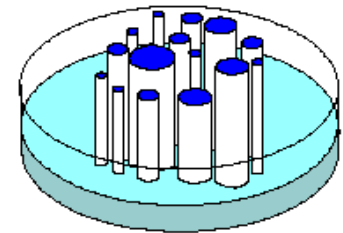
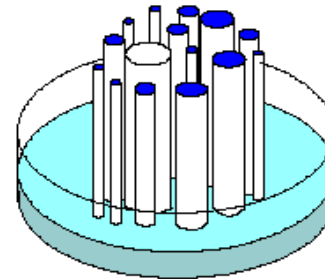
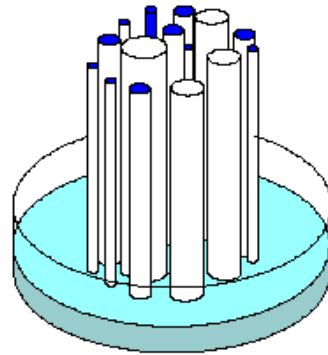
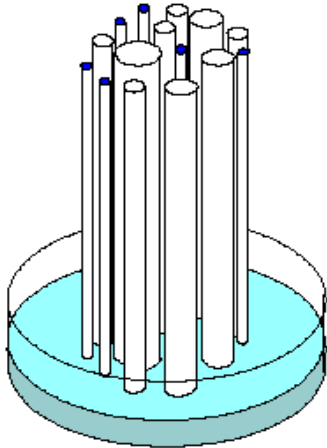
• Evaporation method

..... One-step outflow method (median with 20th and 80th percentiles)

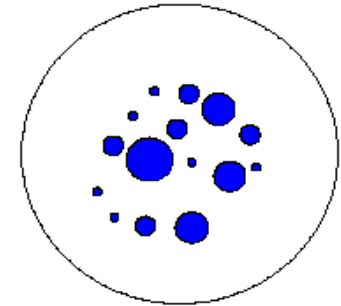
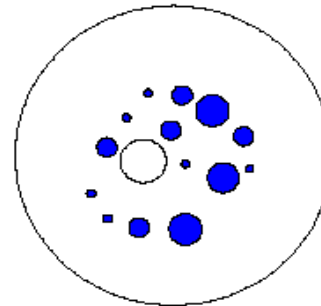
----- Hot-air method (median with minimum and maximum)

Stolte et al. (1994), SSSAJ

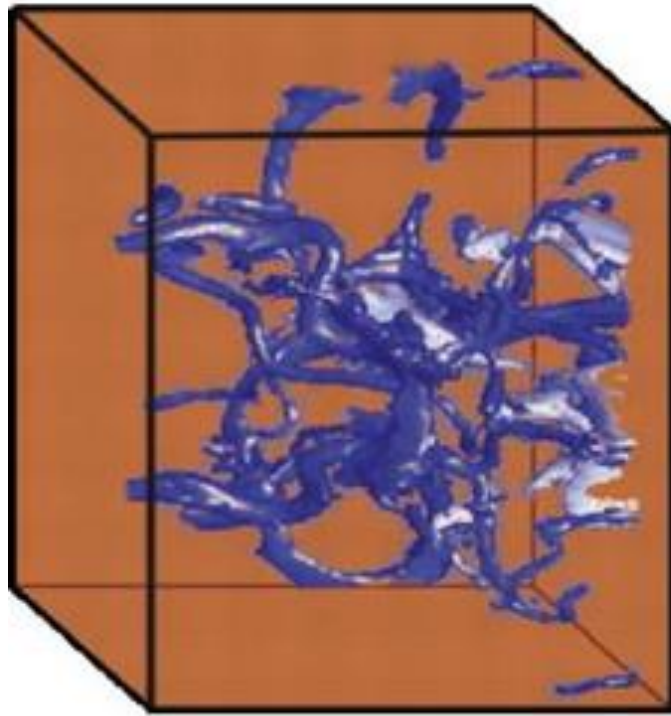
capillary tube model



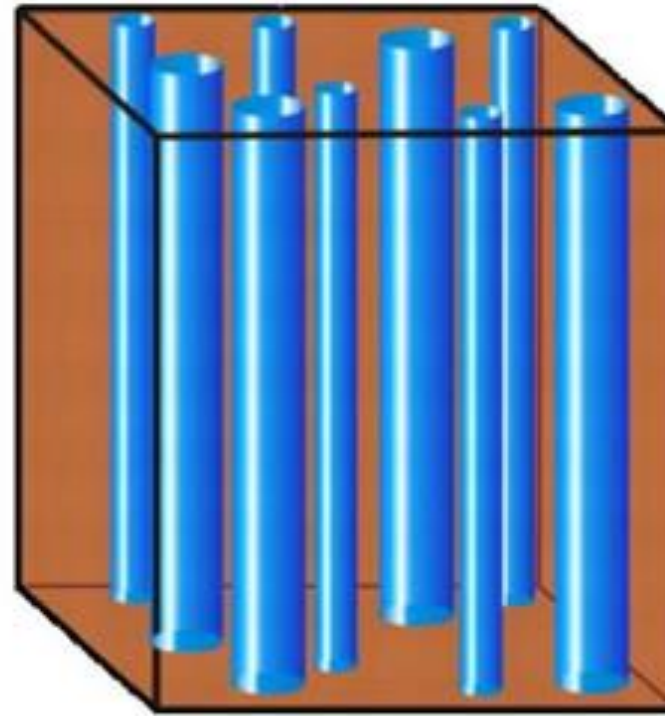
$$K(\theta) = a \cdot \int_0^{\theta} r^2 d\theta$$
$$= b \cdot \int_0^{\theta} \frac{1}{h^2} d\theta$$



Burdine, 1950



Real Media



Idealized Media

Khaleel and Saripalli, VZJ, 2006

Relative conductivity function, $K_r(\theta)$

$$K_r(\theta) = S_e^l \left[\int_0^\theta \frac{1}{h^2} d\theta \bigg/ \int_0^{\theta_s} \frac{1}{h^2} d\theta \right]$$

Tortuosity coefficient l

$$S_e = (\theta - \theta_r) / (\theta_s - \theta_r)$$

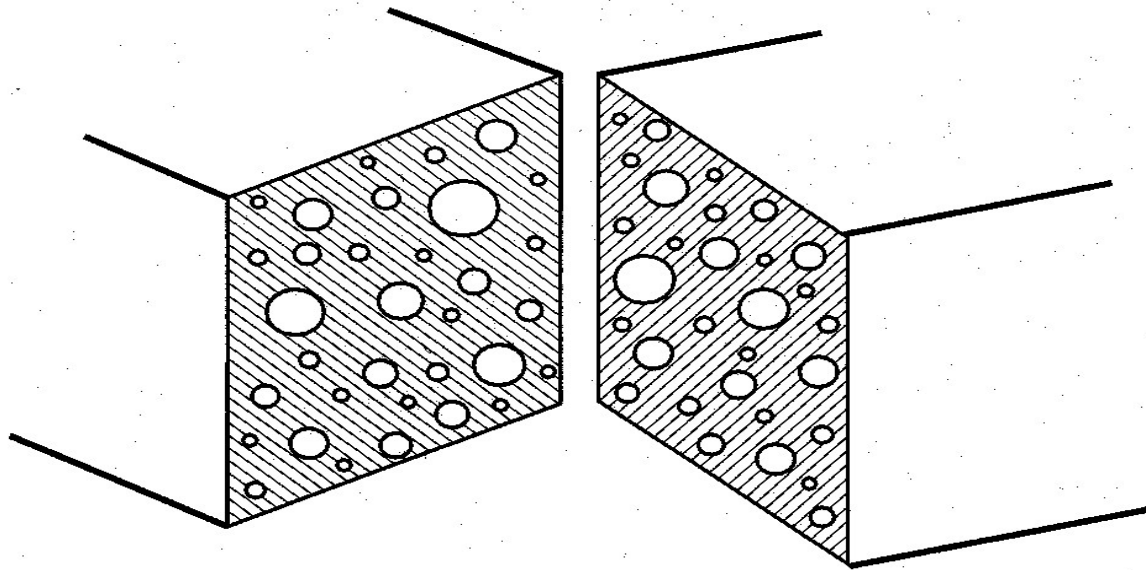
UMS KSAT: Saturated hydraulic conductivity



$$K(\theta) = K_r(\theta) \cdot K_s$$

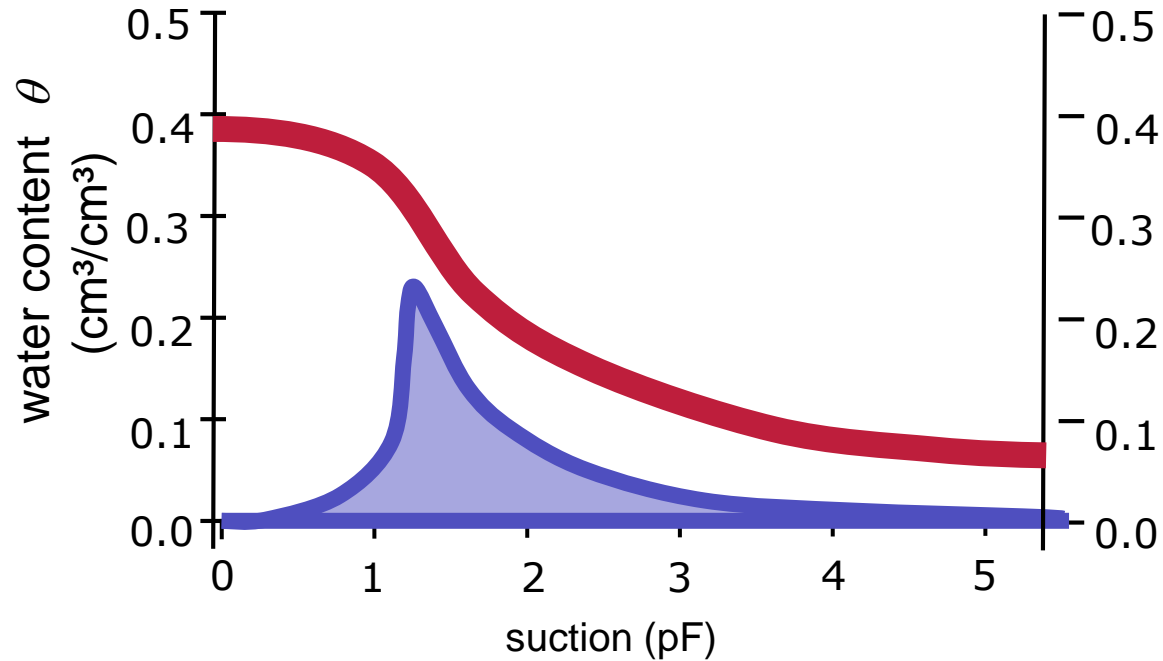
capillary bundle models

cut-and-random-rejoin



Childs and Collis-George, 1950

Capillary bundle model: Mualem (1976)



$$K_r(\theta) = S_e^{0.5} \left[\frac{\int_0^\theta \frac{1}{h} d\theta}{\int_0^{\theta_s} \frac{1}{h} d\theta} \right]^2$$

van Genuchten-Mualem model (1980)

retention function

$$S_e = \left(\frac{1}{1 + |\alpha h|^n} \right)^m$$

conductivity function

$$K = K_s S_e^l \left[1 - \left(1 - S_e^{1/m} \right)^m \right]^2$$

α [m^{-1}], n [-], m [-]

$S_e(h)$ [-]

K_s [m s^{-1}]

l [-]

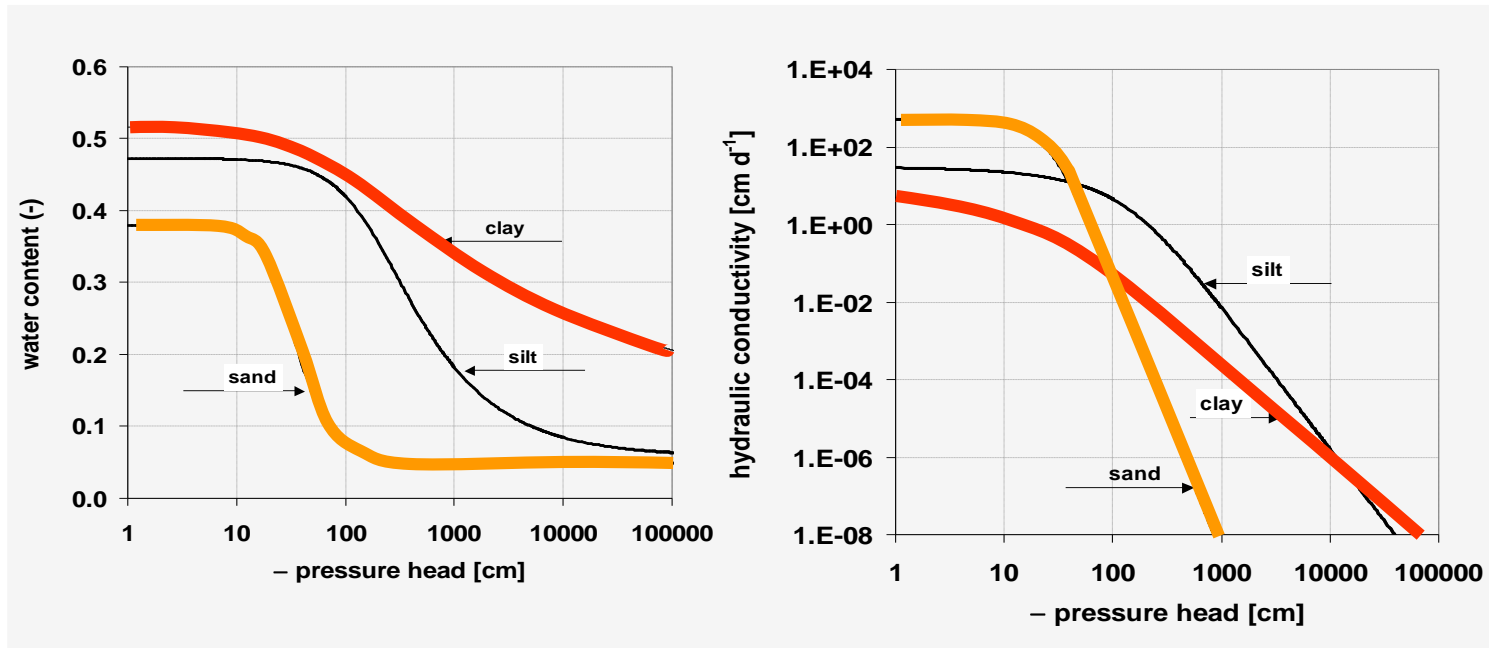
van Genuchten parameters

effective saturation

saturated conductivity

tortuosity coefficient

van Genuchten-Mualem parametrization



Typical hydraulic properties of differently textures soils Left: Water retention curves.
Right: Unsaturated conductivity curves (Durner und Flühler, 2005).

So far: nothing new.

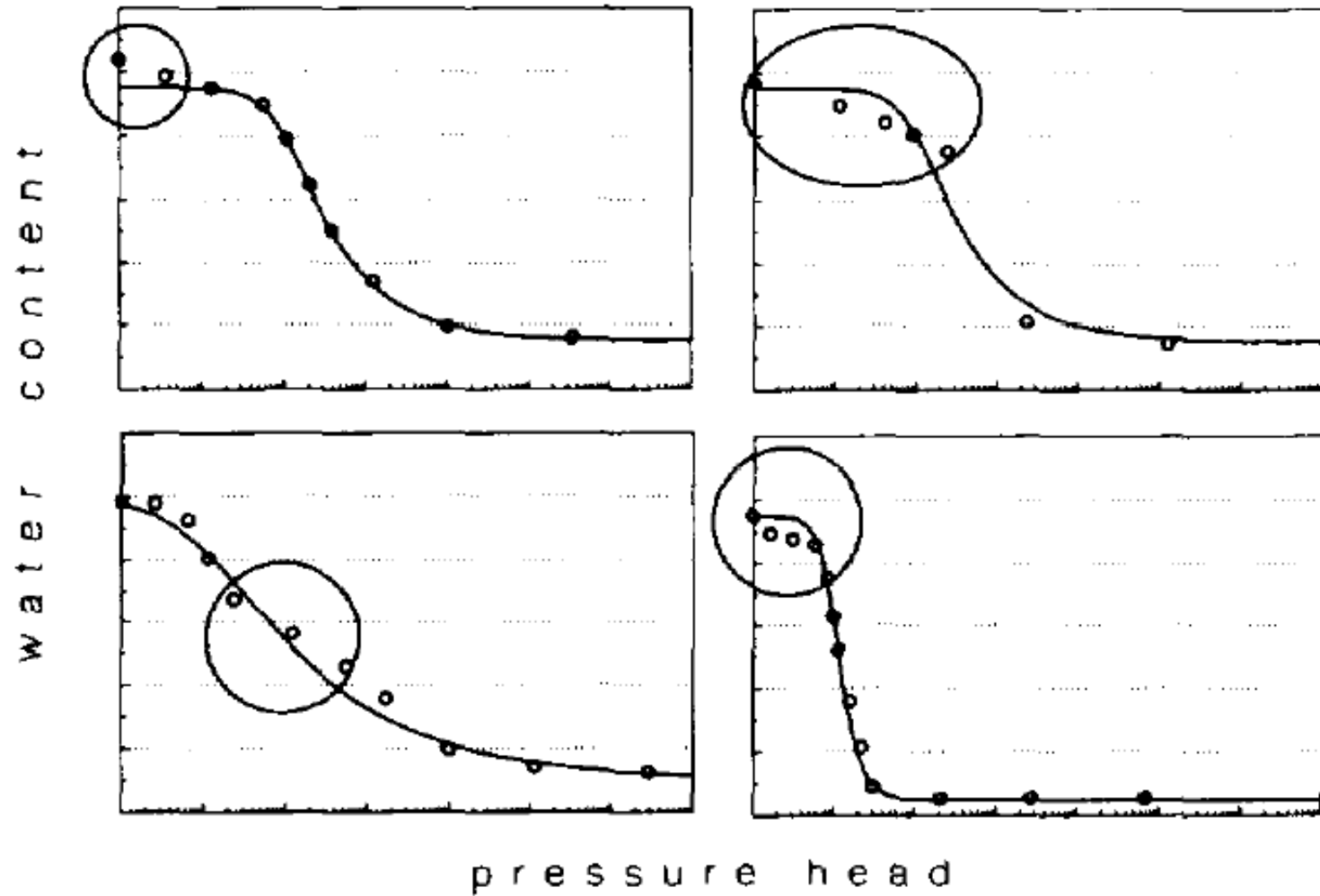
Part 3:

A closer look on hydraulic properties



The wet range

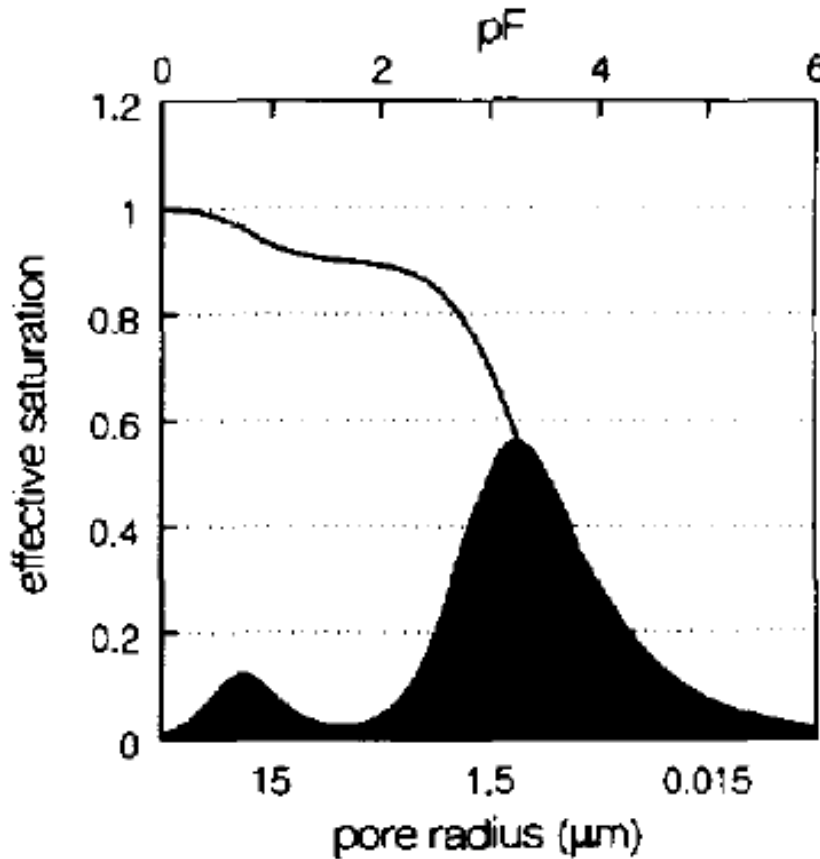
Wet range: small changes ...



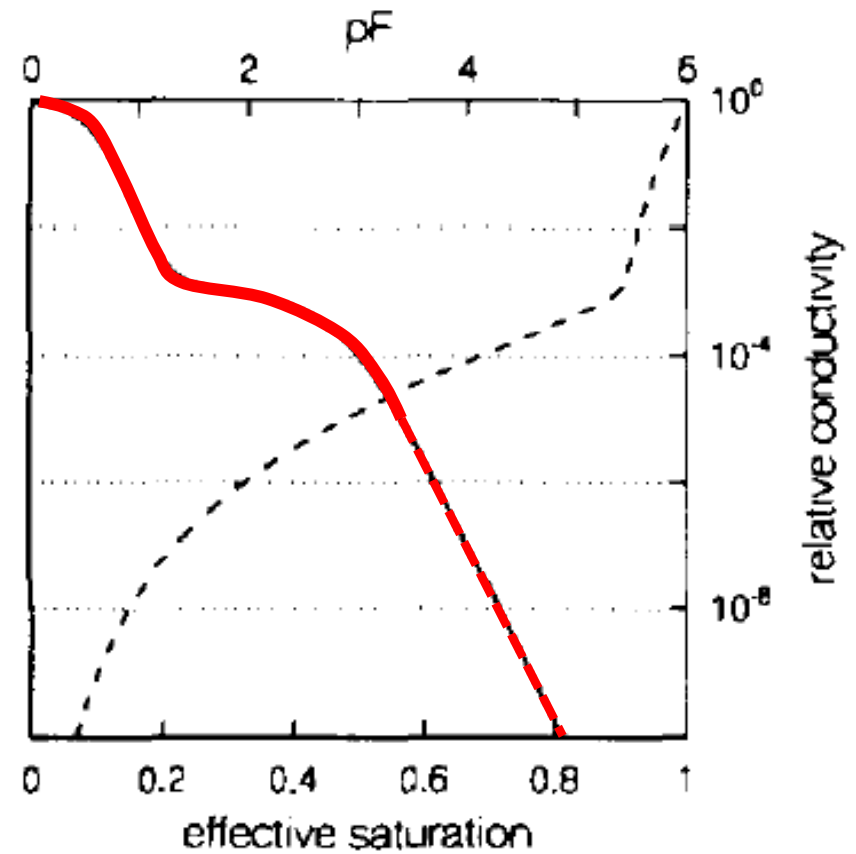
Durner (1994), WRR

Wet range: small changeshave big effects

Water retention function



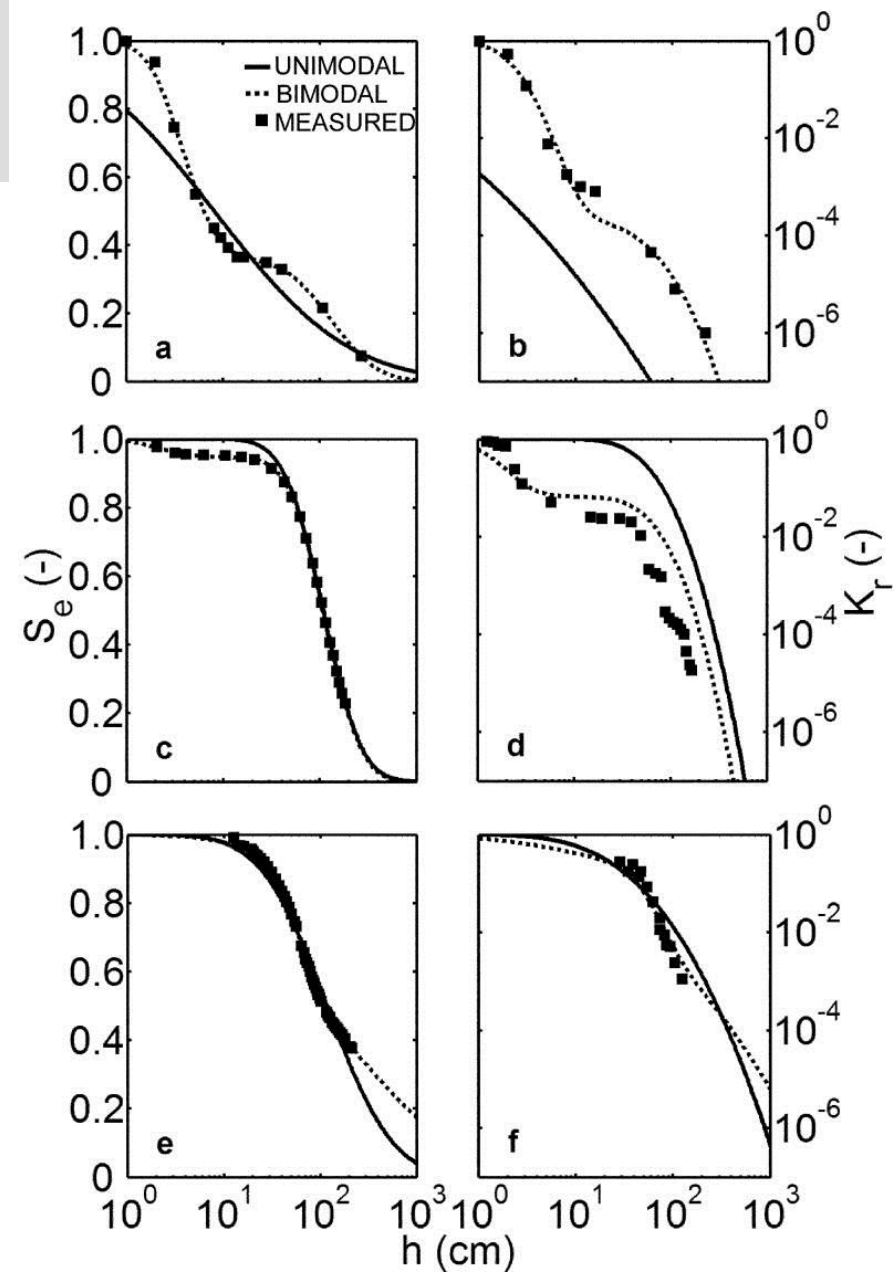
Hydraulic conductivity function



Durner (1994), WRR

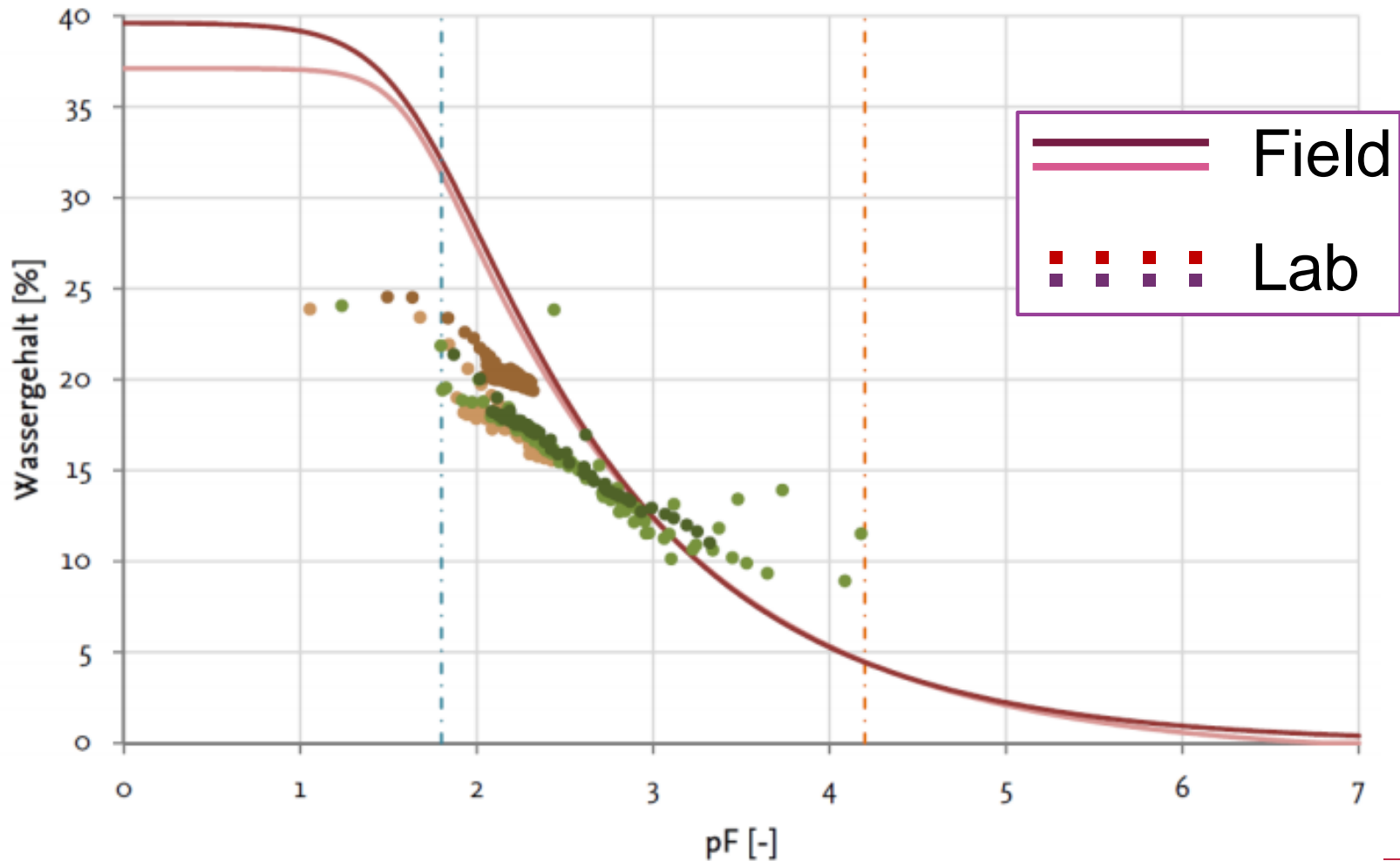
Using Bimodal Lognormal Functions to Describe Soil Hydraulic Properties

N. Romano et al., SSSAJ, 2010



The mid range

Mid range: fundamentally wrong perception?

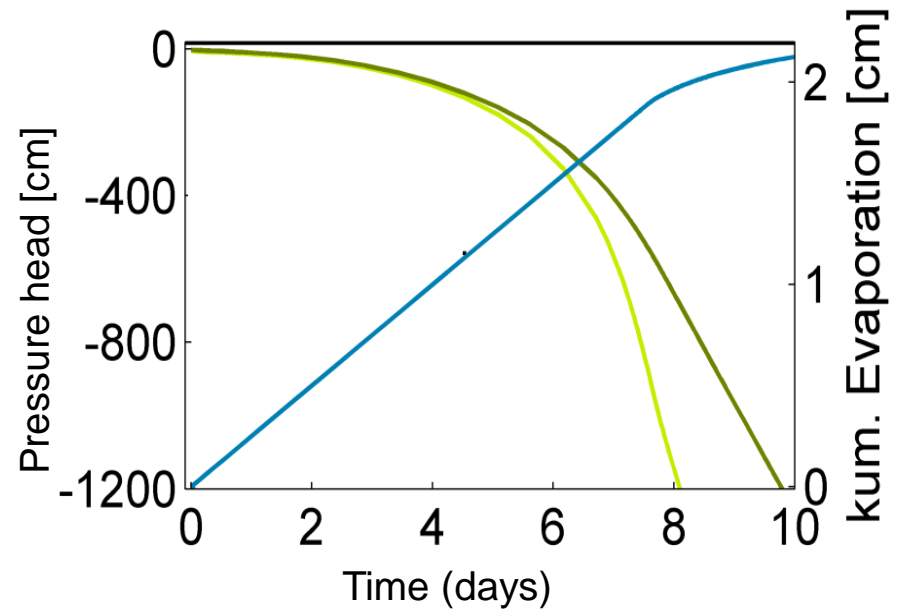
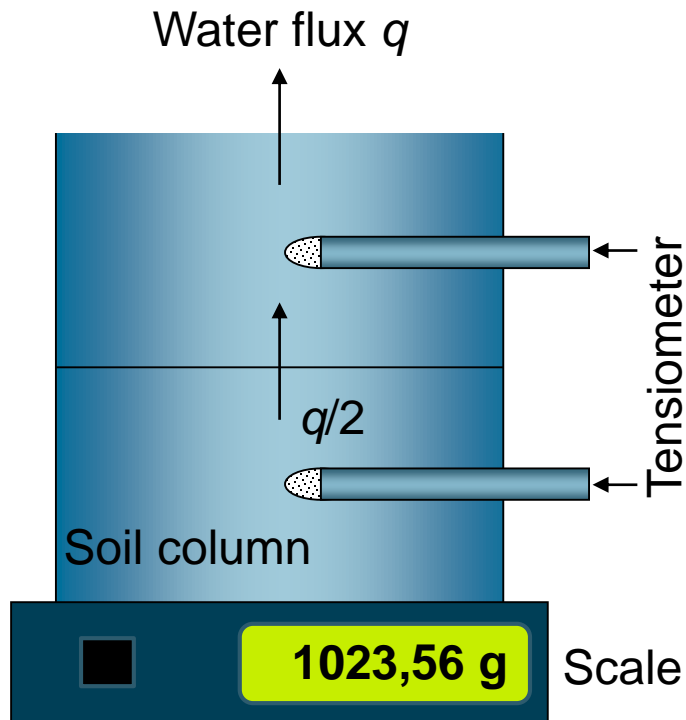


Hydraulic Properties: Challenges

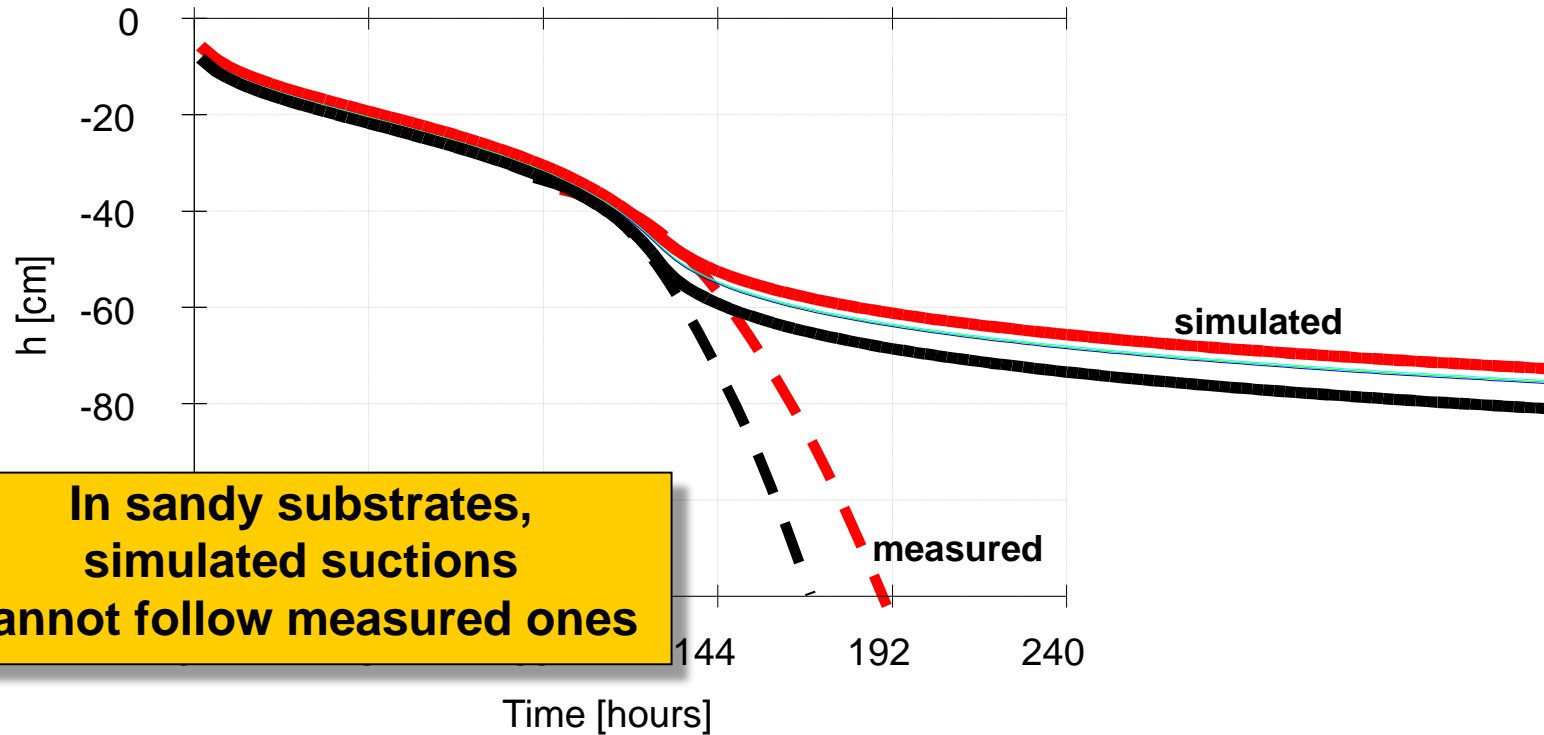
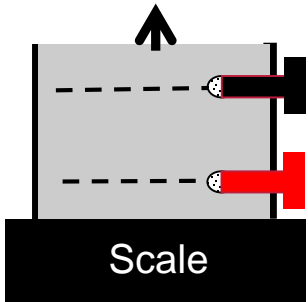
- Hysteresis and dynamic effects
- Temporally varying hydrophobicity (mucilage, biofilms, OM)
- Shrinking and swelling porous media
- Heterogeneity and Upscaling
- Chemical and biological feedbacks
- T-dependence and thermohydraulic coupling

The dry range

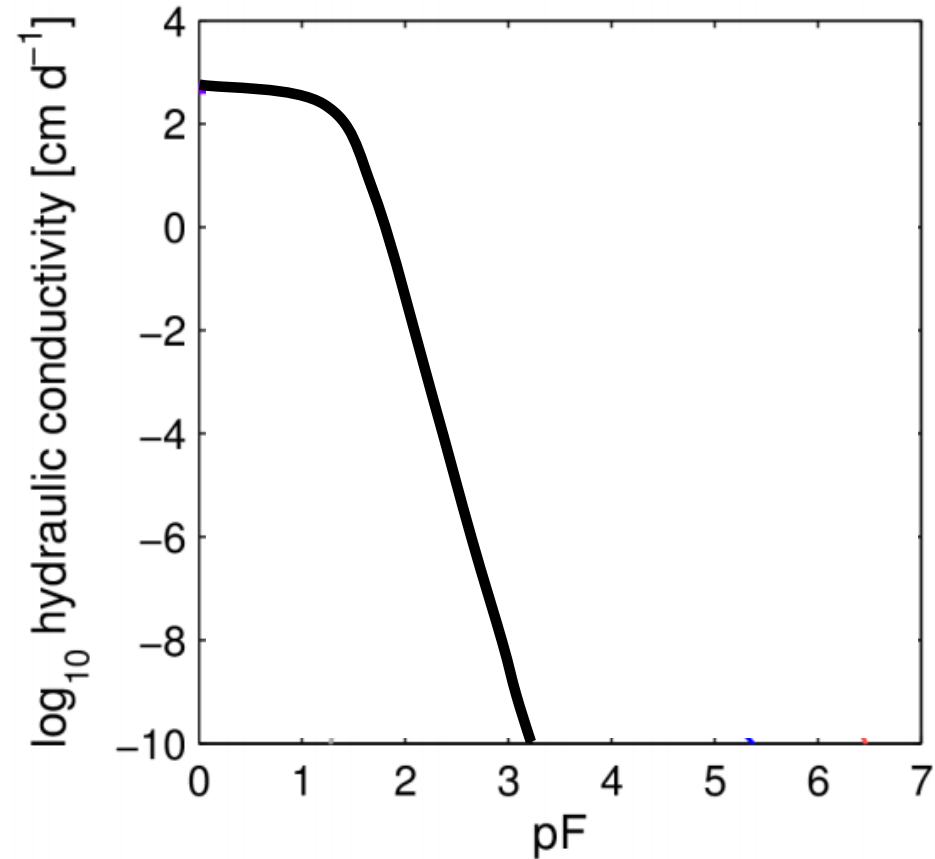
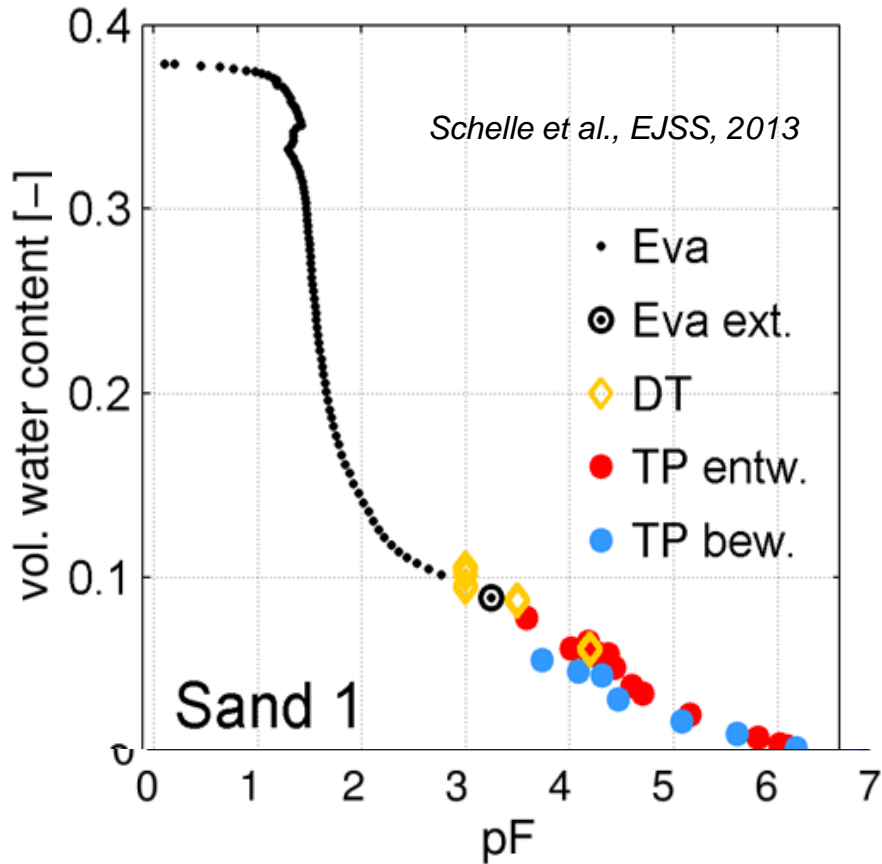
Dry end: Evaporation Experiments



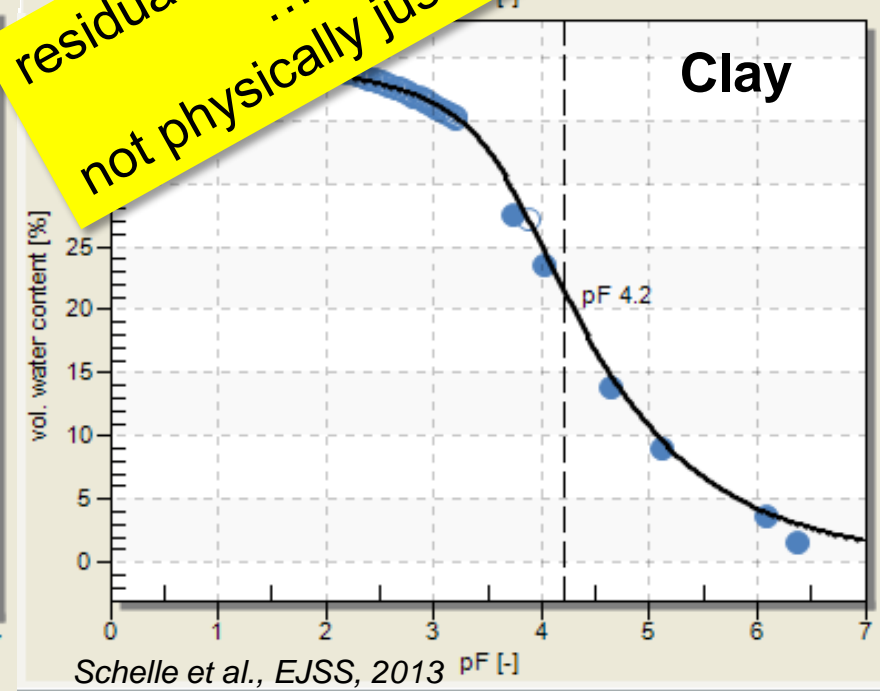
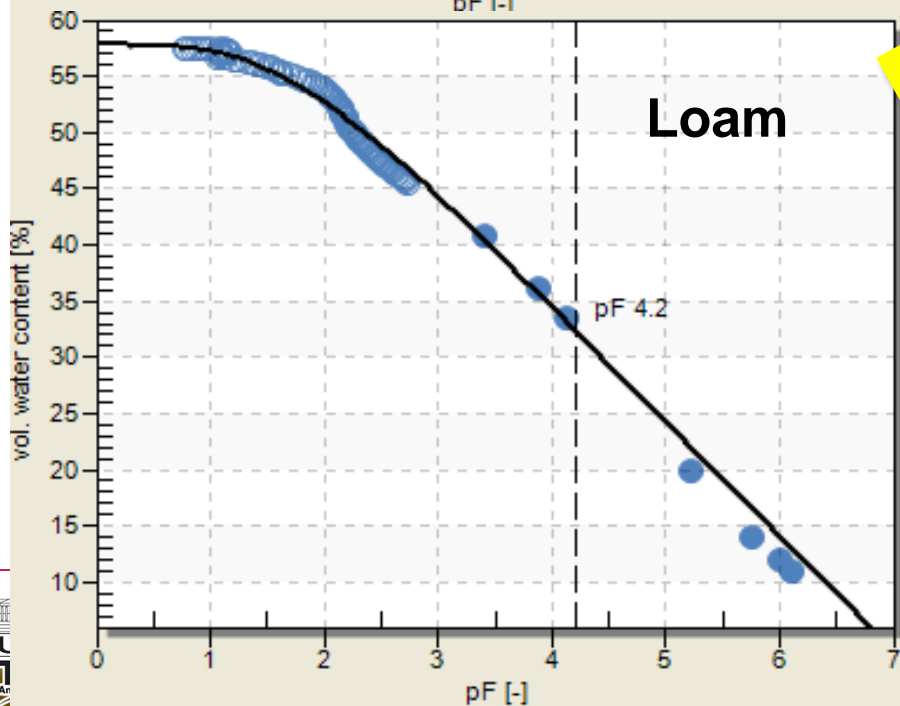
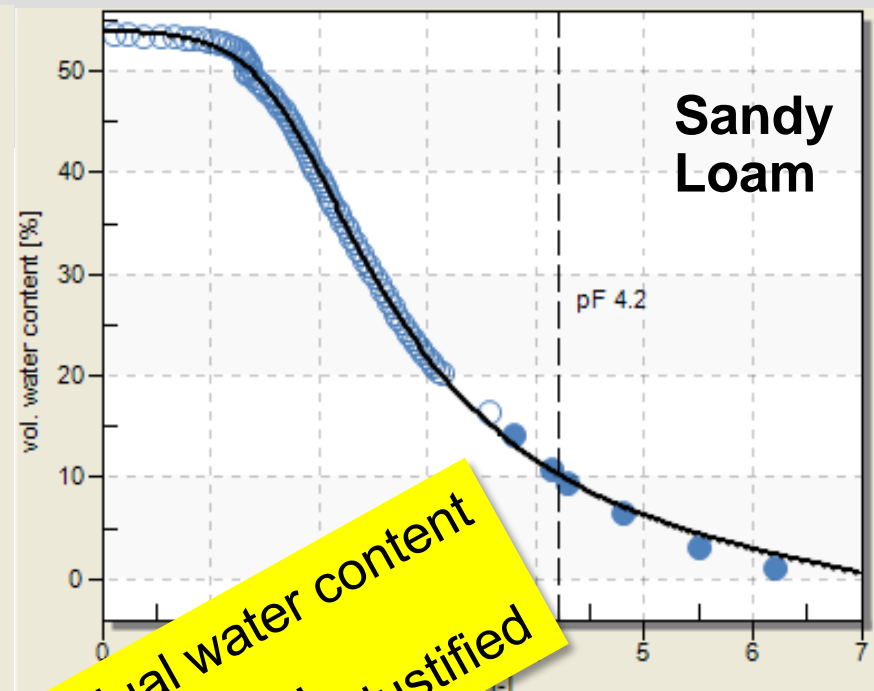
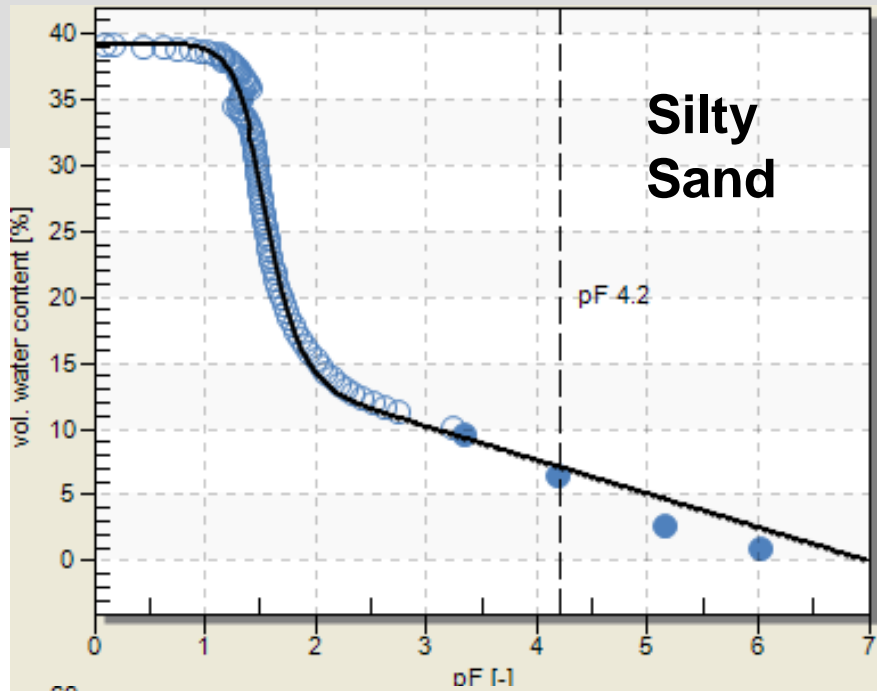
Observation (1)



Observation (2)

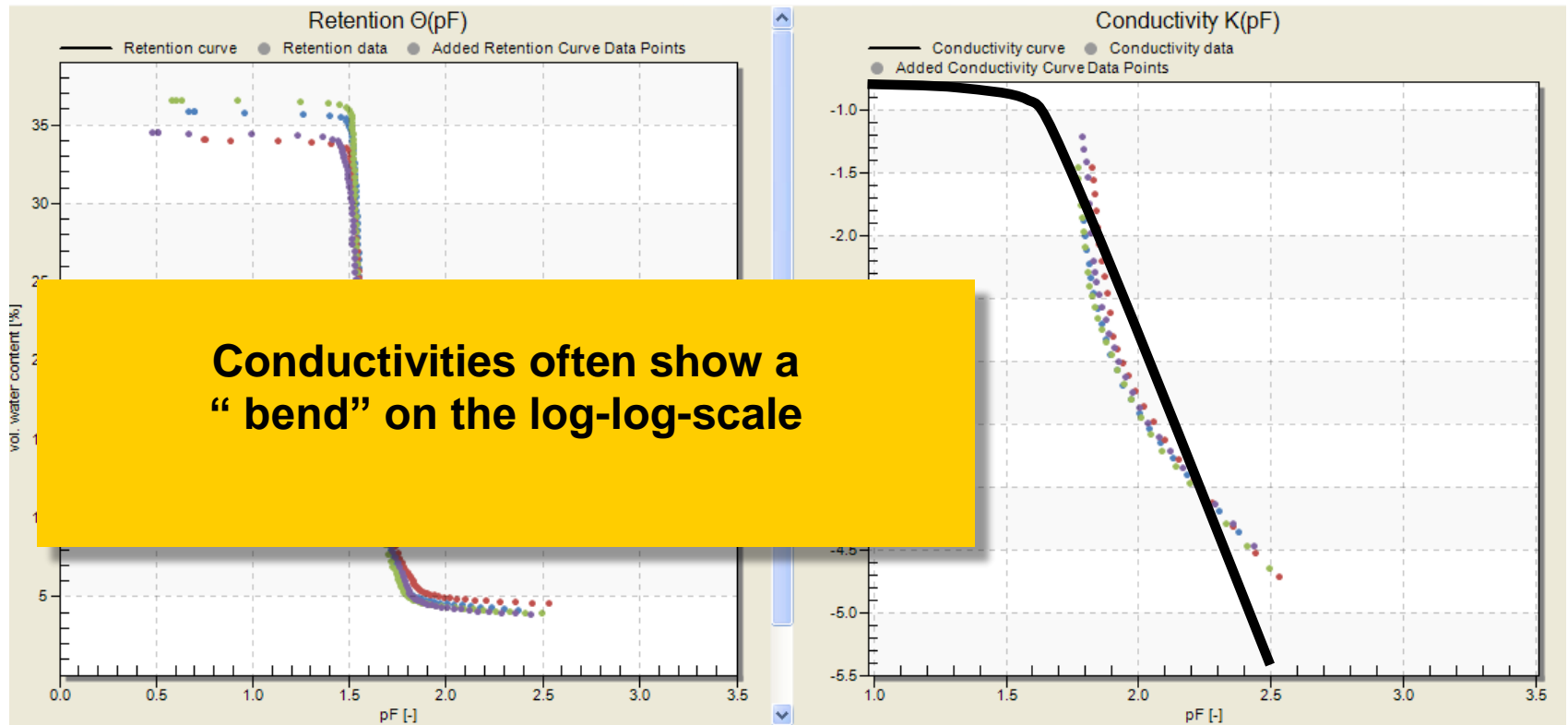


e.g. Nimo, 1991



residual water content
...
not physically justified

Observation (3)

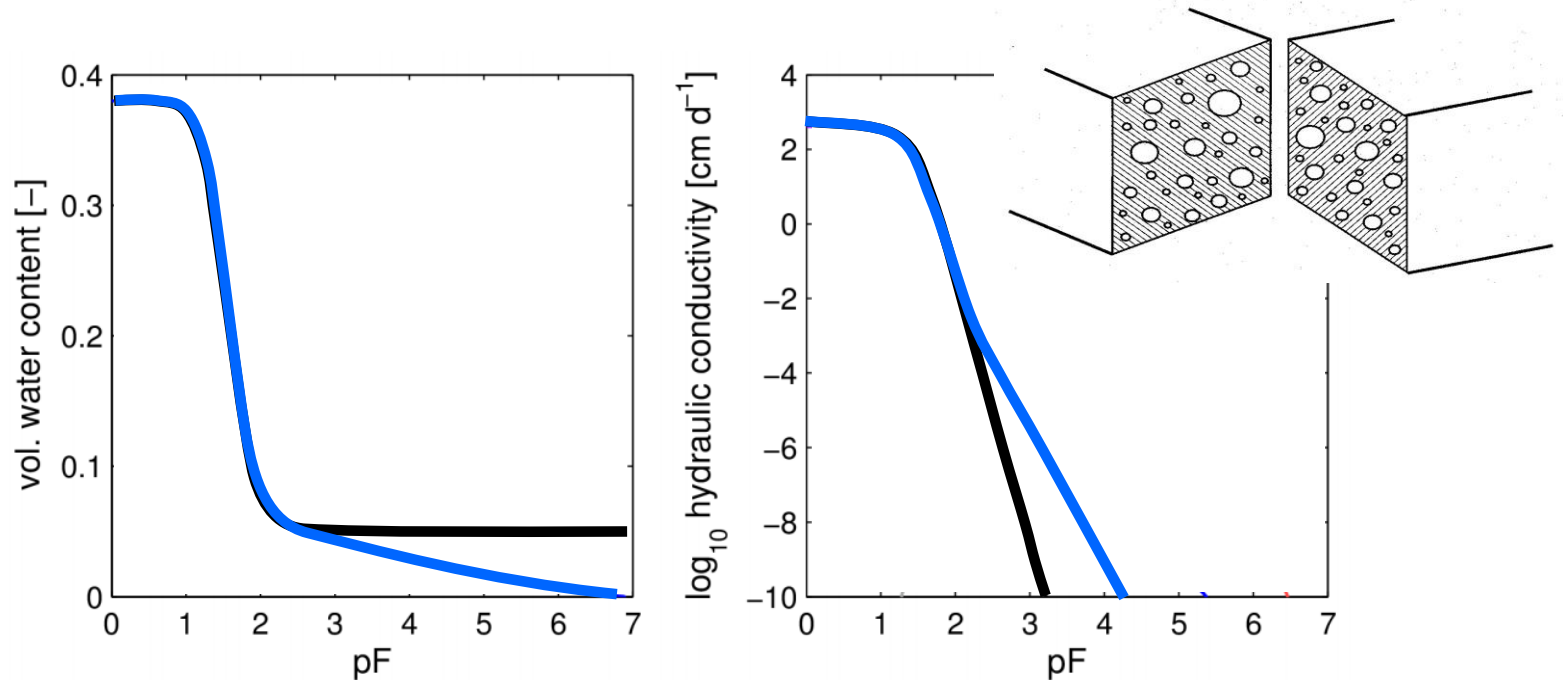


Conductivities often show a
"bend" on the log-log-scale

Part 4: Better shaped hydraulic properties

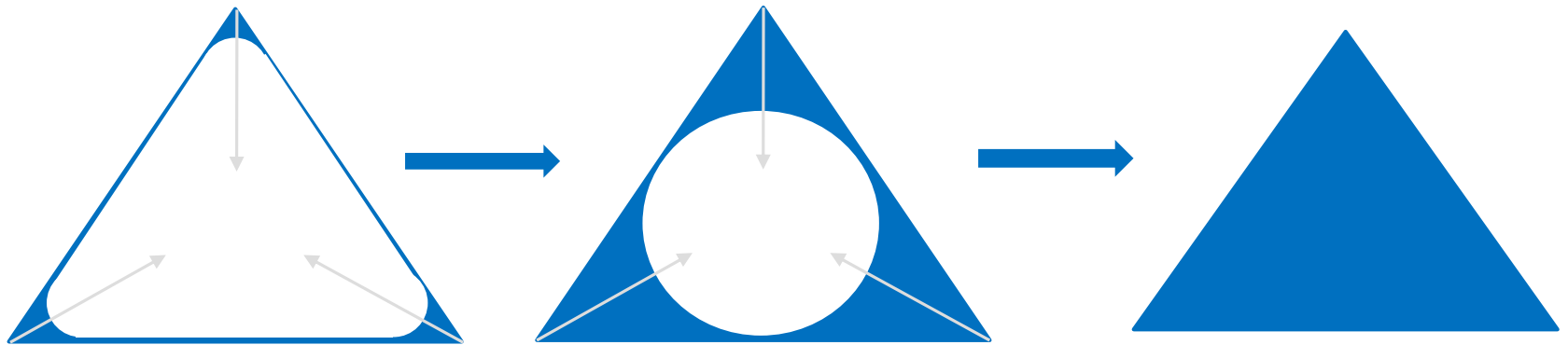


Better SHP ...



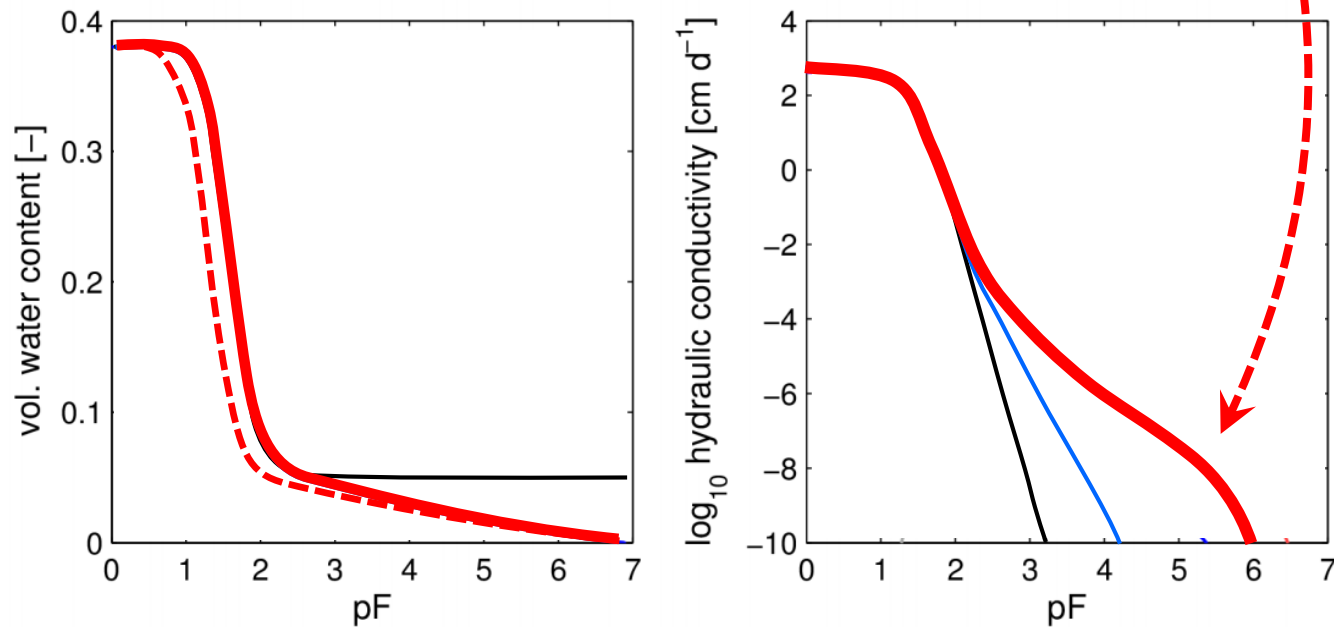
Ross et al. 1991 | Campbell-Shiozawa 1992 | Rossi and Nimmo 1994 | Fredlund-Xing 1994 | Fayer-Simmons 1995 | Morel-Seytoux and Nimmo 1999 | Webb 2002 | Groenevelt and Grant 2004 | Khlosi et al 2006 | ...

Adding a film-flow + corner-flow component



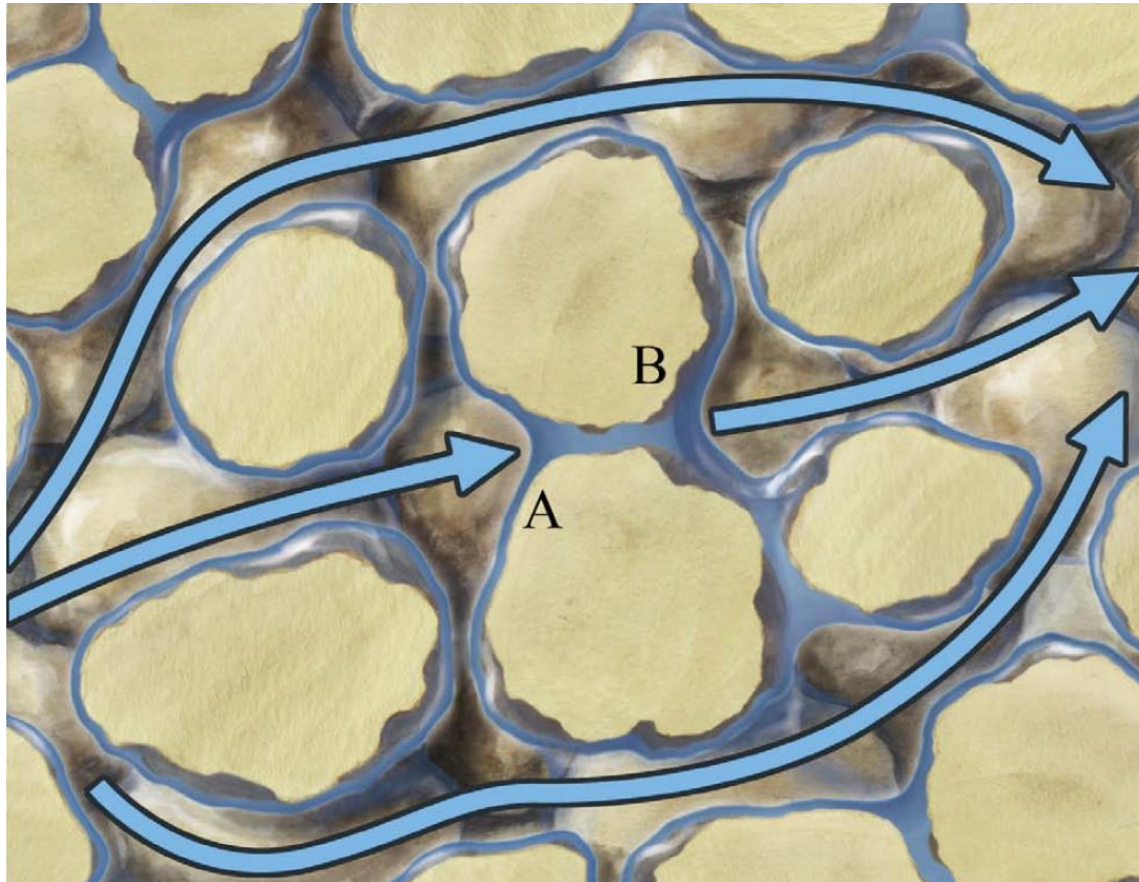
Tuller-Or 2001

... even better SHP ...



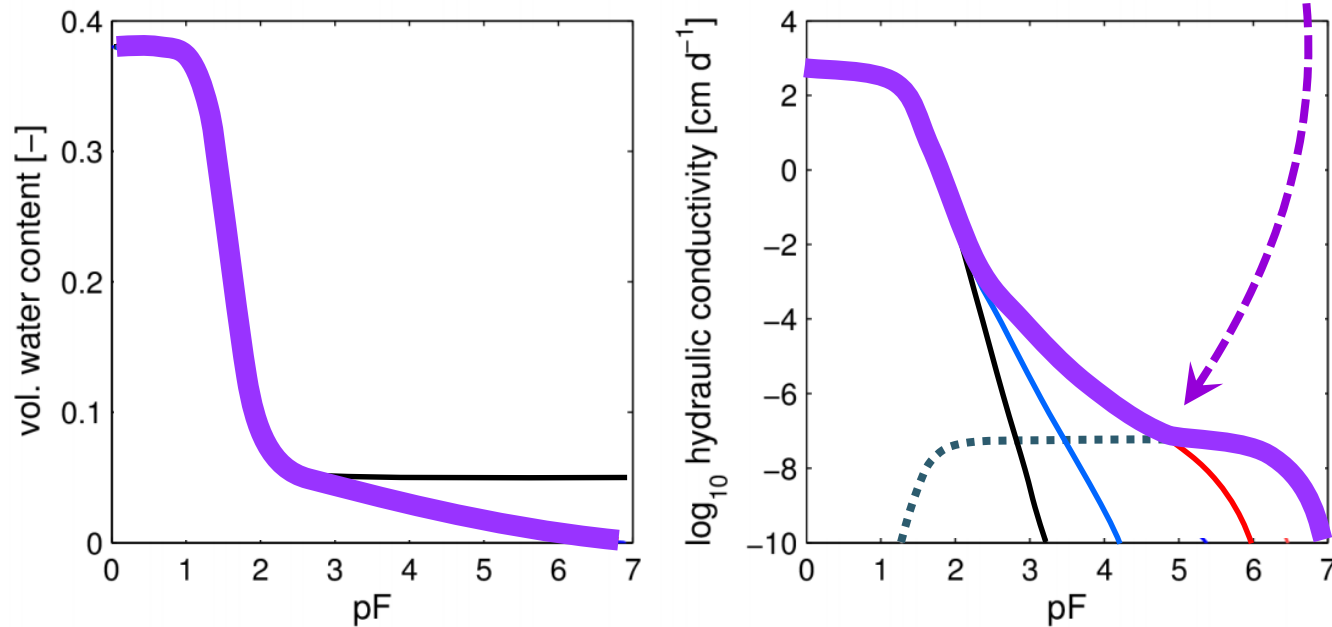
Tuller-Or 2001 | Peters-Durner 2008 | Tokunaga 2009 | Lebeau-Konrad 2010 | Zhang 2011 | Diamantopoulos-Durner, 2013 | ...

(Enhanced) diffusion of water vapor



(Enhanced) diffusion of water vapor by condensation and evaporation in presence of isolated liquid water
(Fig. by Shahraneen und Or, 2012)

Really good SHP !



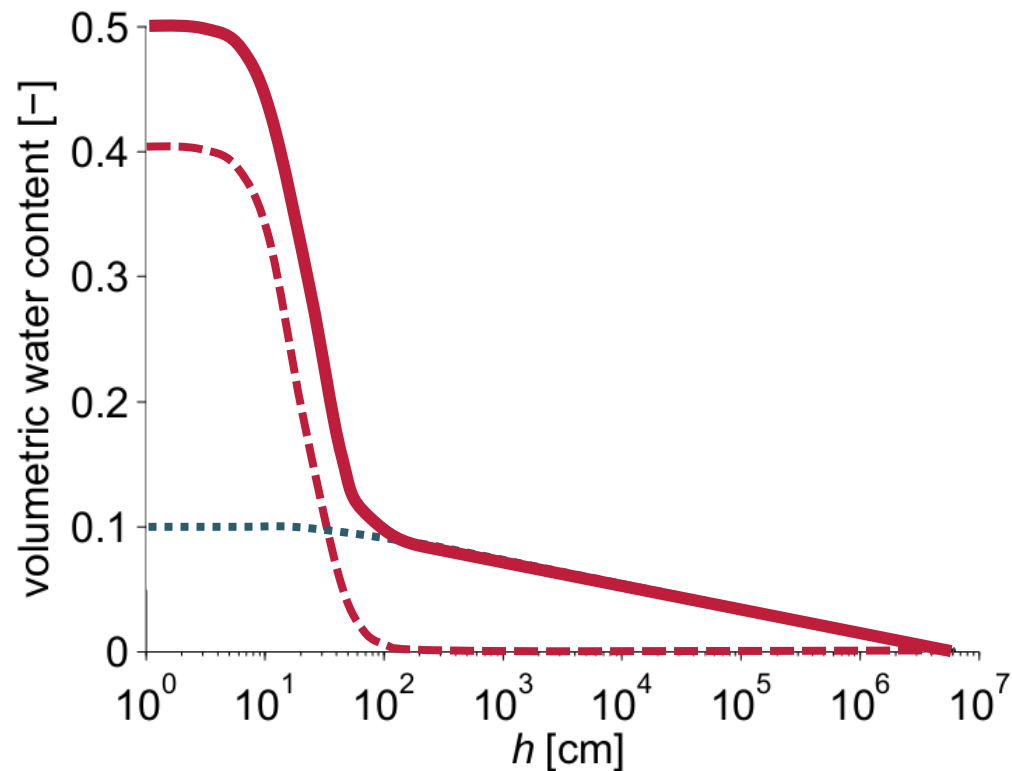
Philip and de Vries 1957 | Saito et al. 2006

A parsimonious parameterization: The PDI model



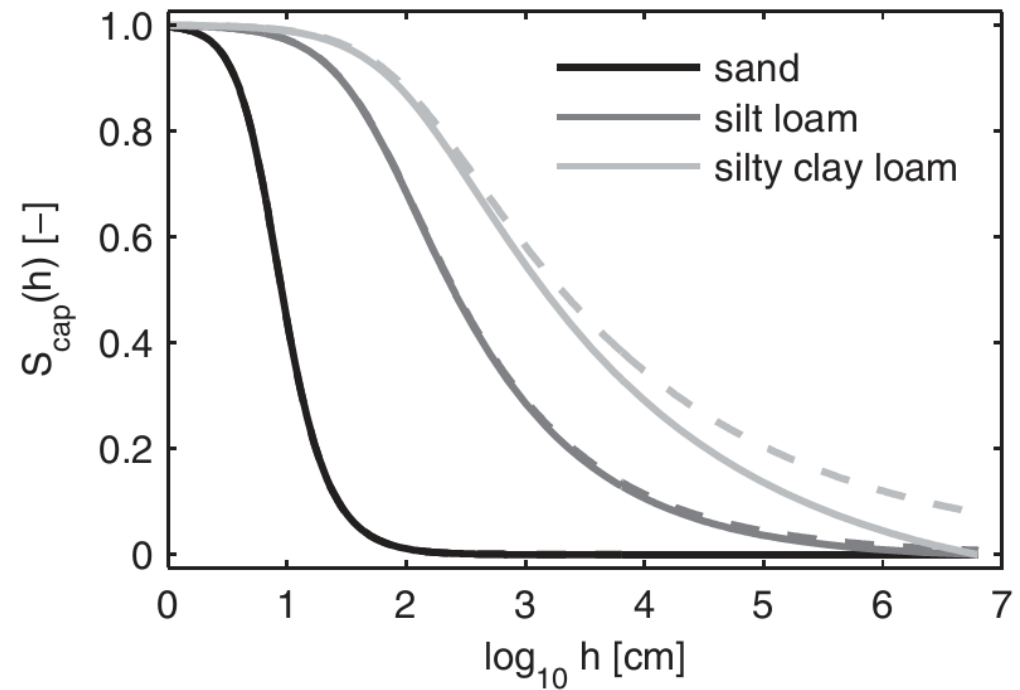
PDI model: retention curve (Peters 2013 | Iden and Durner 2014 | Peters 2014)

$$\theta(h) = (\theta_s - \theta_r) S_{\text{cap}}(h) + \theta_r S_{\text{ad}}(h)$$



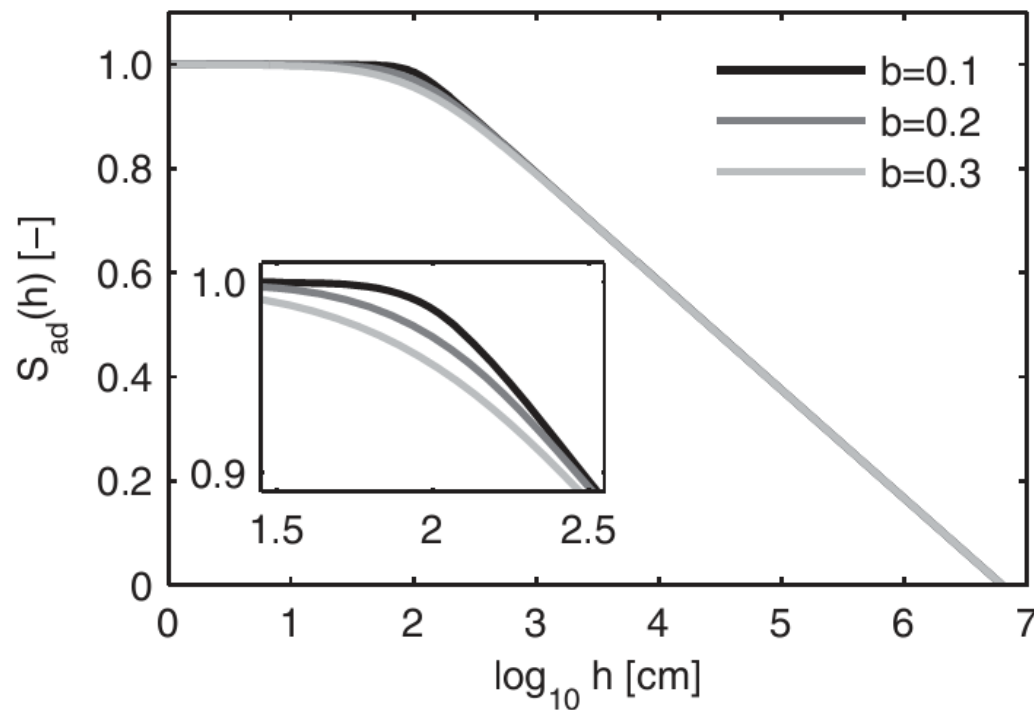
PDI model: capillary part

$$S_{\text{cap}}(h) = \frac{\Gamma(h) - \Gamma(h_0)}{1 - \Gamma(h_0)}$$



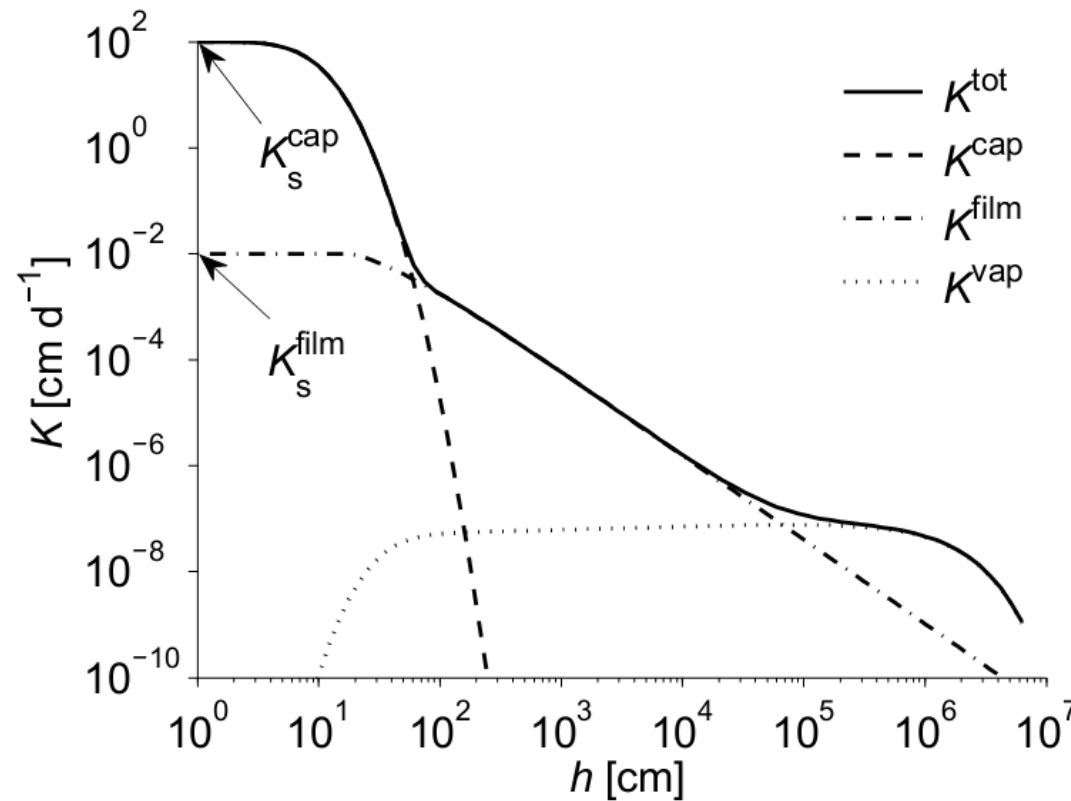
PDI model: non-capillary part

$$S_{ad}(x) = 1 + \frac{1}{x_a - x_0} \left\{ x - x_a + b \ln \left[1 + \exp \left(\frac{x_a - x}{b} \right) \right] \right\}$$

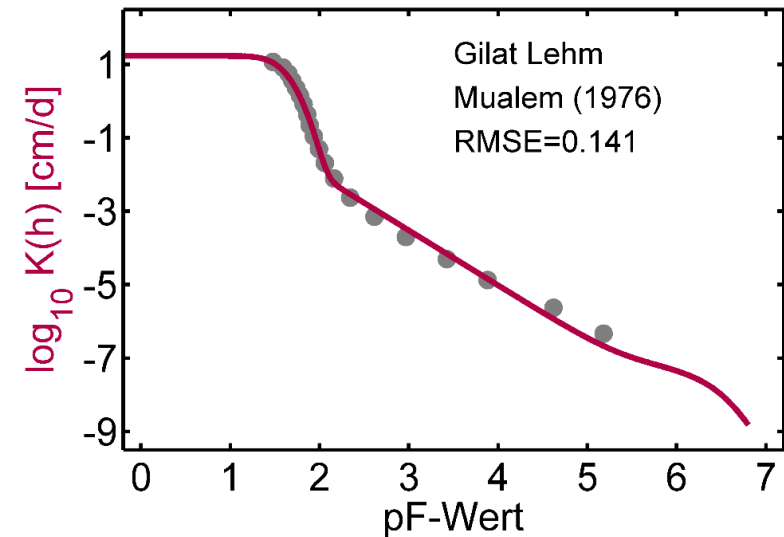
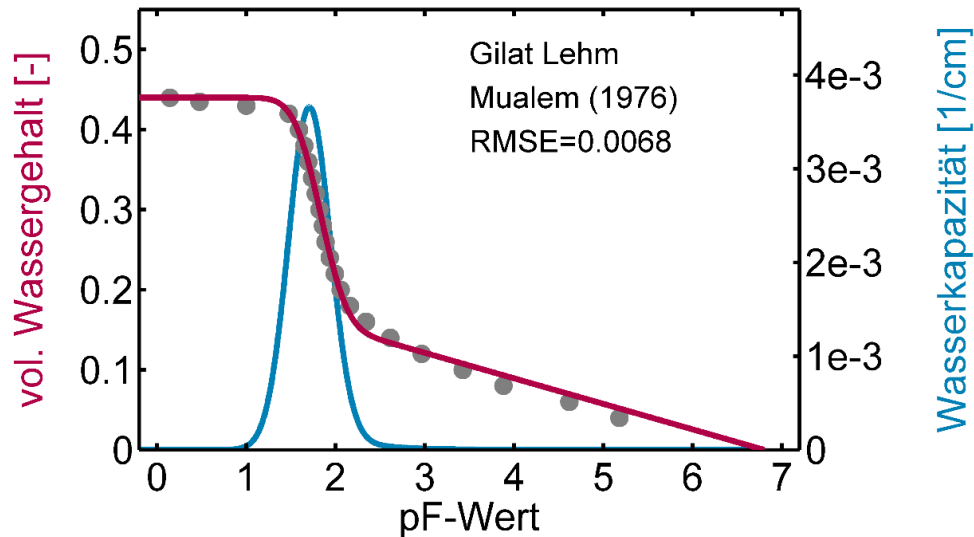
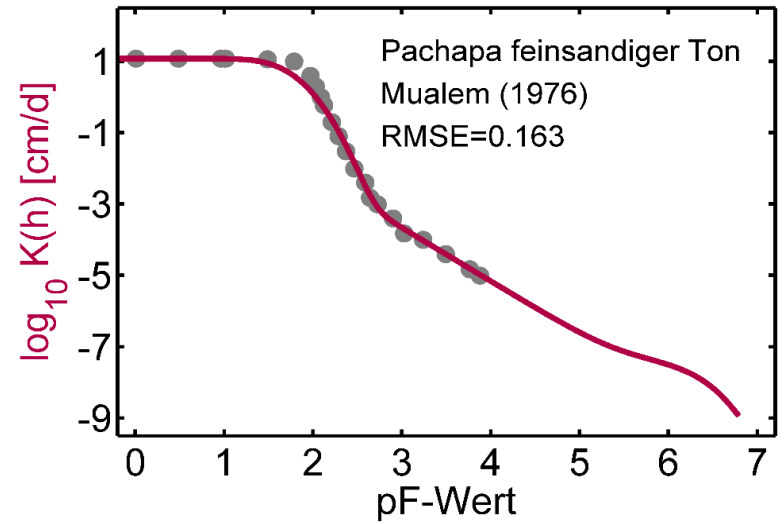
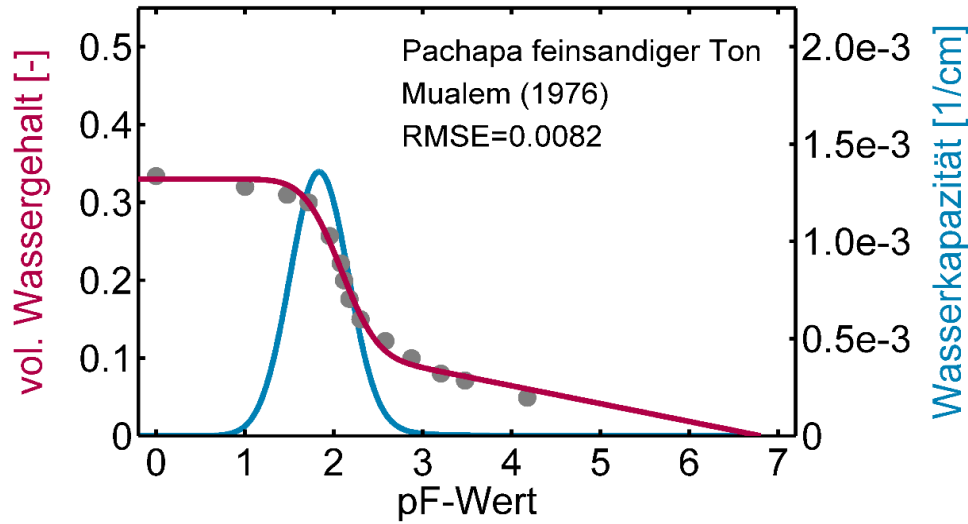


PDI model: conductivity curve

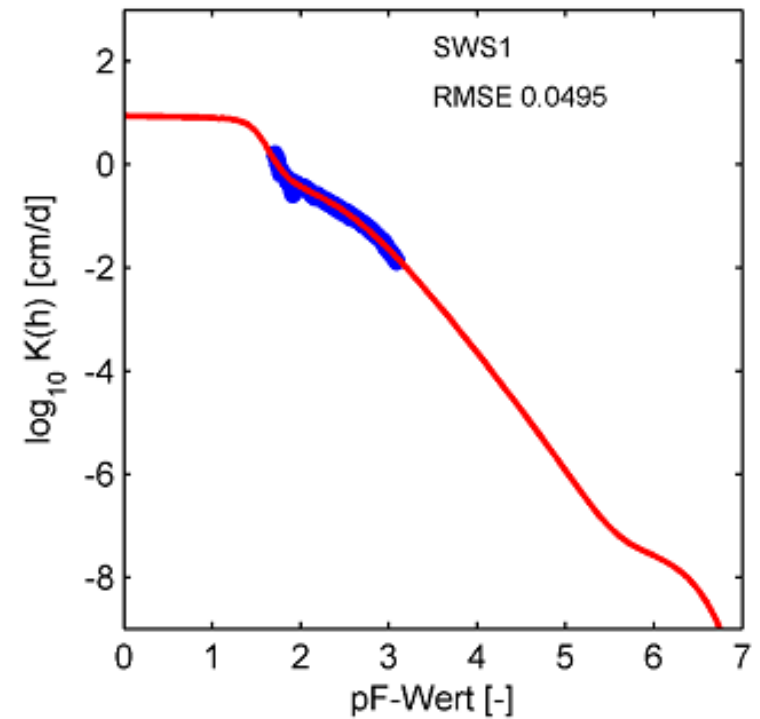
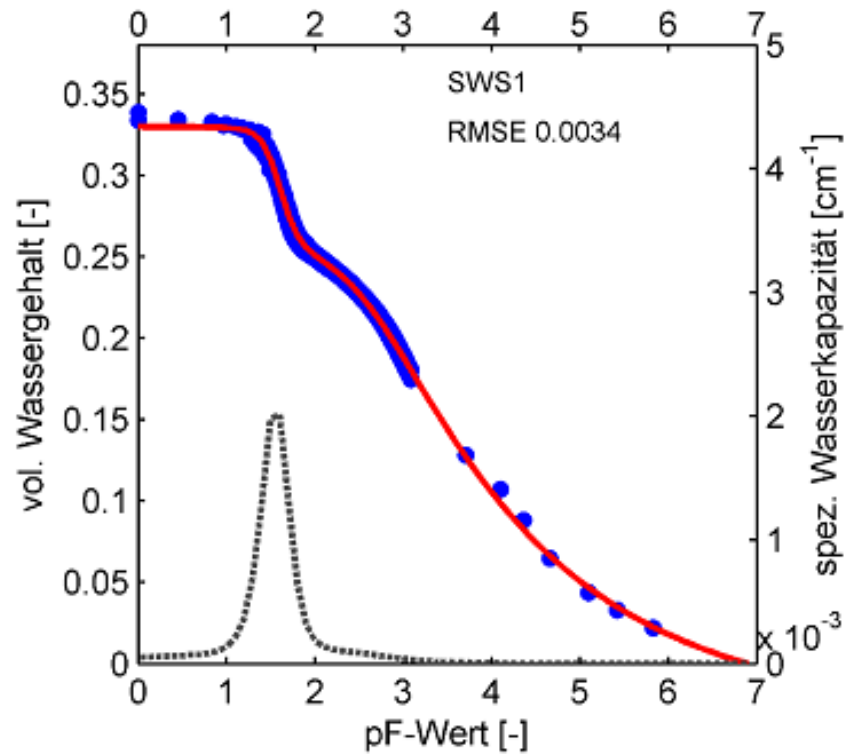
$$K^{lv} = K^{liq} + K^{vap} = K^{cap} + K^{film} + K^{vap}$$



PDI model fitted to data



PDI model fitted to data



Summary: Properties of the PDI model

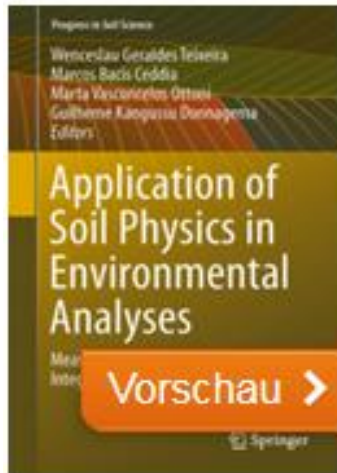
- Always reaches zero water content at oven dryness
- Continuous water capacity function
- RETC: no more parameters than traditional models, e.g., $\alpha, n, \theta_s, \theta_r$
- Very robust
- Implemented in the free HYPROP-FIT software

Information Fitting Export

Soil hydraulic model selection

	original	PDI	bimodal	bimodal PDI
Brooks-Corey	<input type="radio"/>			
Fredlund-Xing	<input type="radio"/>	<input type="radio"/>		<input type="radio"/>
Kosugi	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>
van Genuchten $m=1-1/n$	<input checked="" type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>
van Genuchten $mnvar$	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>	<input type="radio"/>

Model: VG Model Code: 1100
Source: van Genuchten (1980)
Description: traditional constrained van Genuchten-Mualem model



Chapter 17 Hydraulic Properties and Non-equilibrium Water Flow in Soils

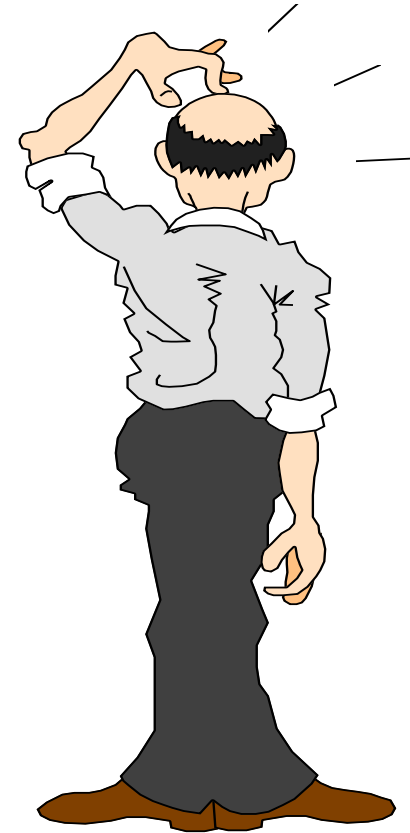
Wolfgang Durner, Efstathios Diamantopoulos,
Sascha C. Iden, and Benedikt Scharnagl

Abstract Accurate knowledge of hydraulic properties for unsaturated soils is critical in the estimation of soil water fluxes by simulation models that are based on the Richards equation. The purpose of this chapter is to review the characteri-

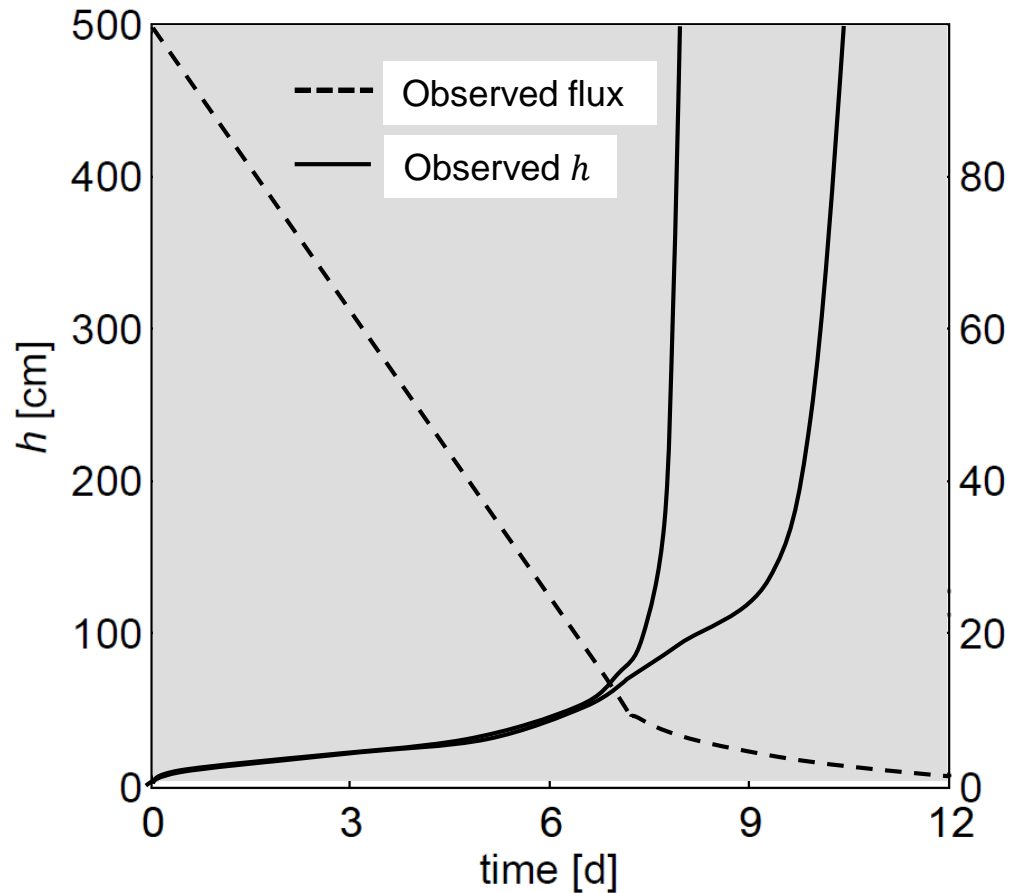
Does that matter ?



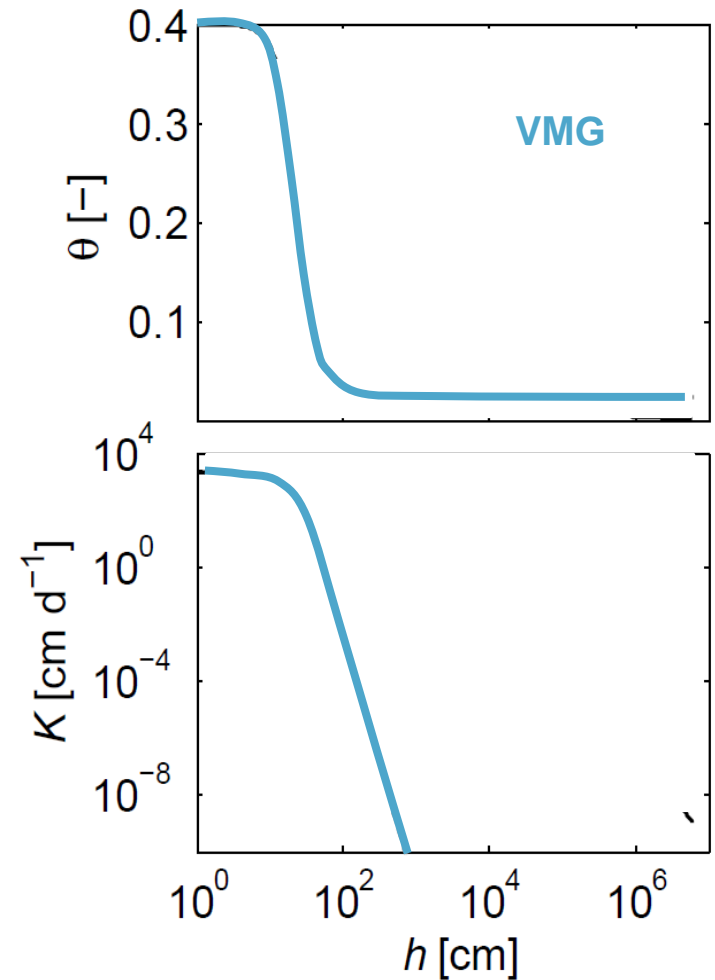
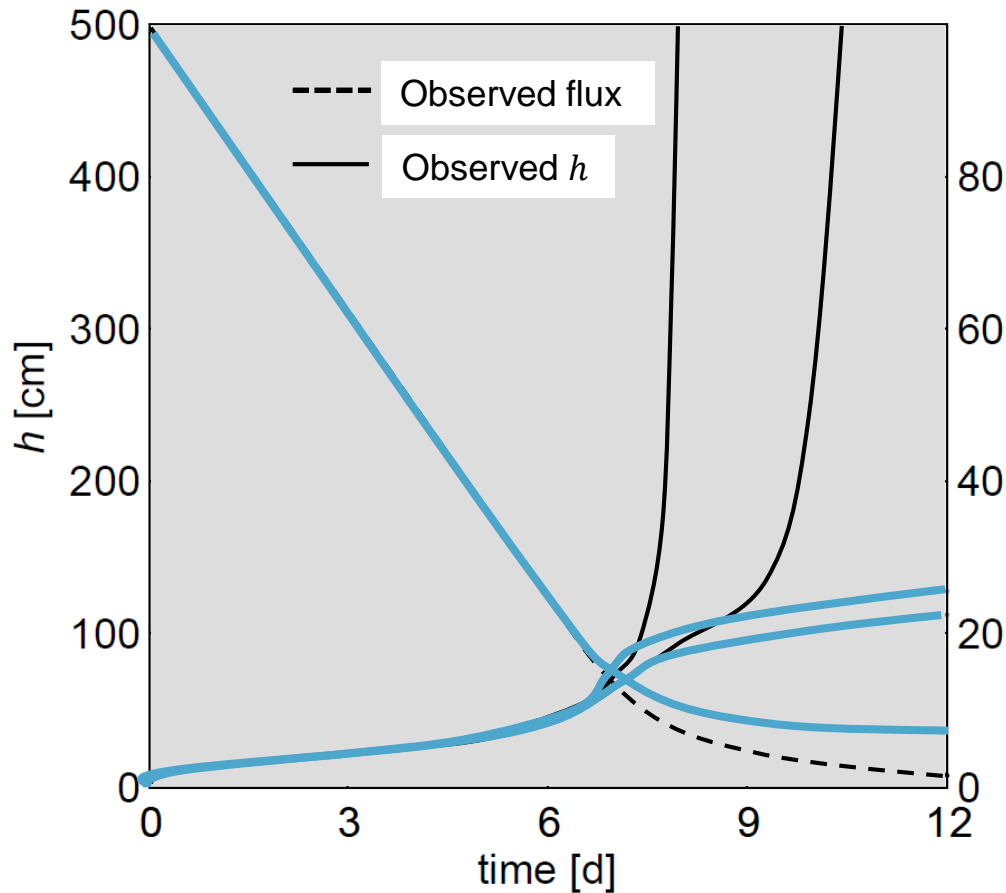
Part 5: Application Example



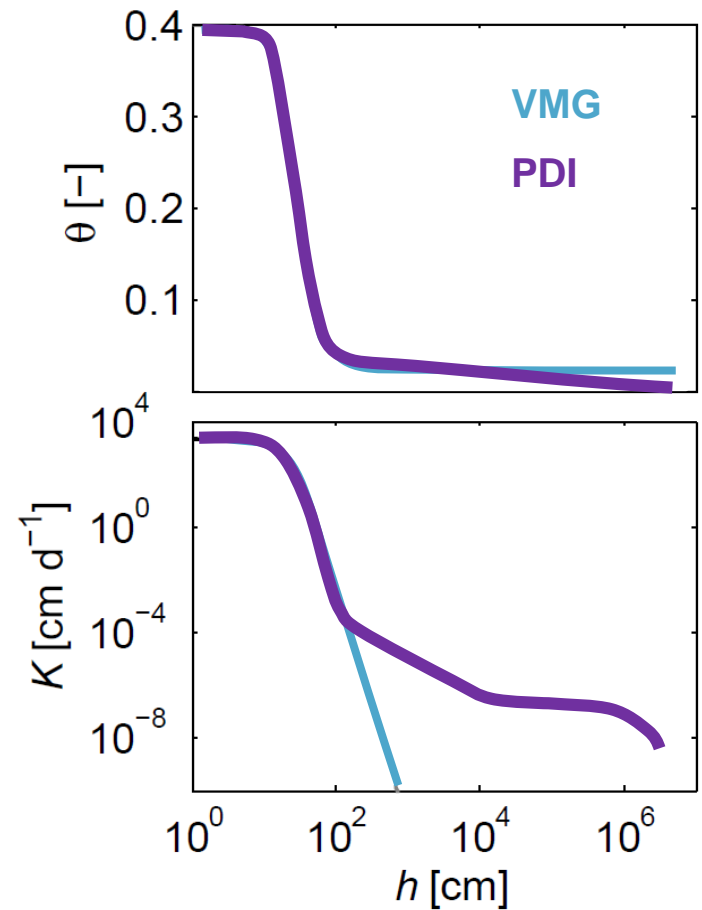
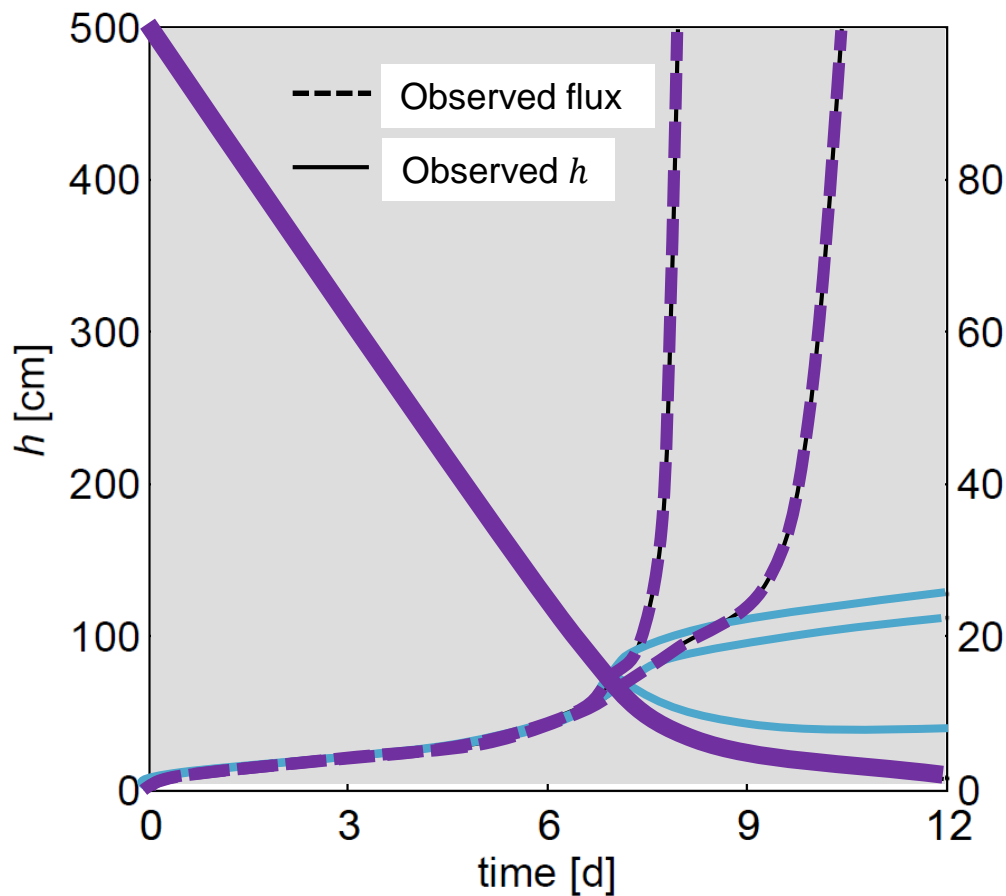
Does that matter (1)?



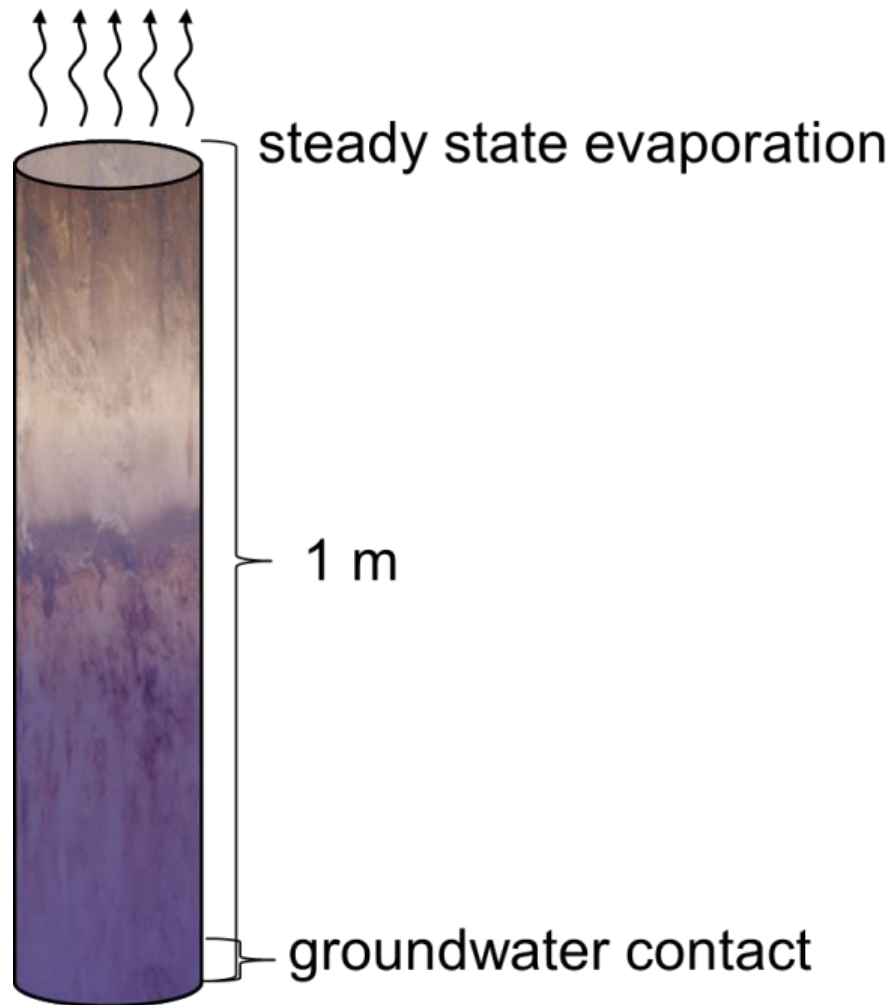
Does that matter (1)?



Does that matter (1)?

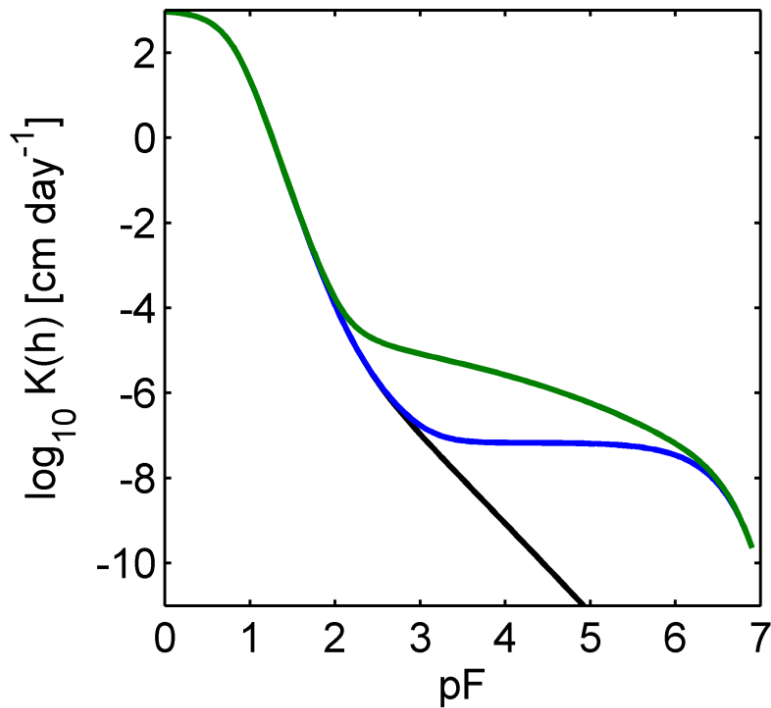


Does that matter (2)?

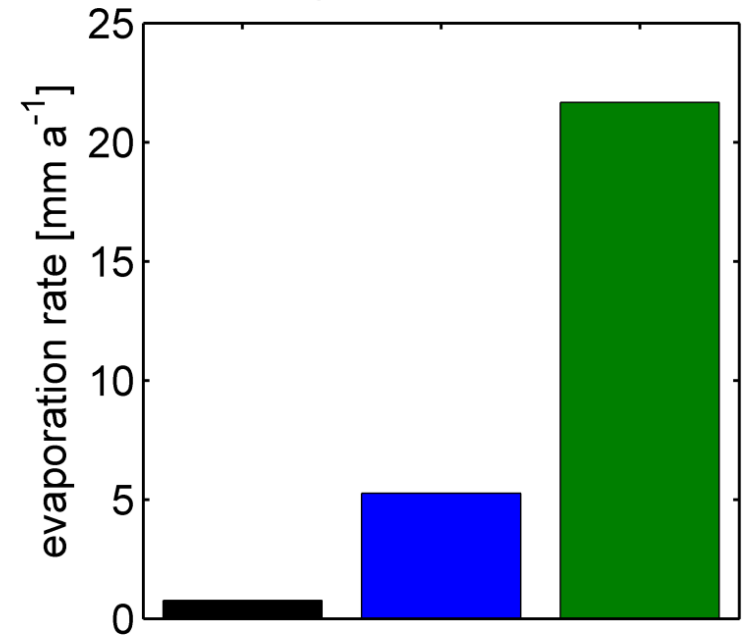


Does that matter (2)?

conductivity function

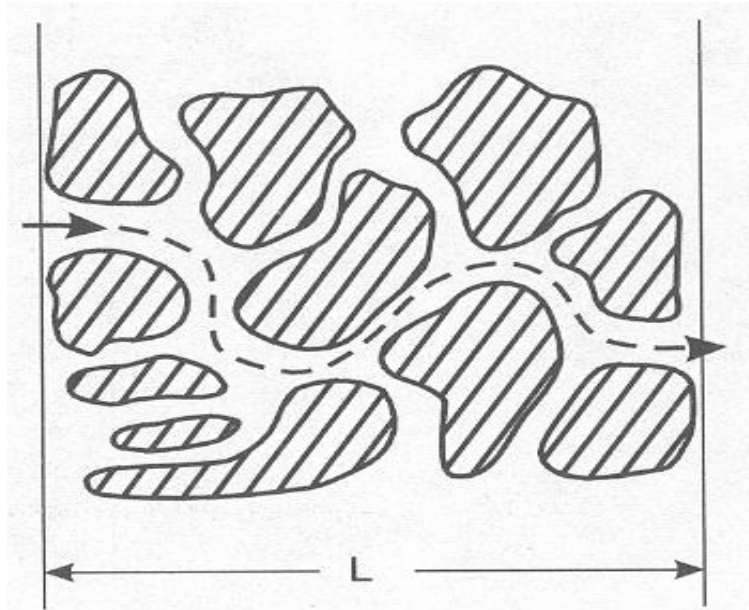


steady-state evaporation



Final Excursus: The tortuosity coefficient

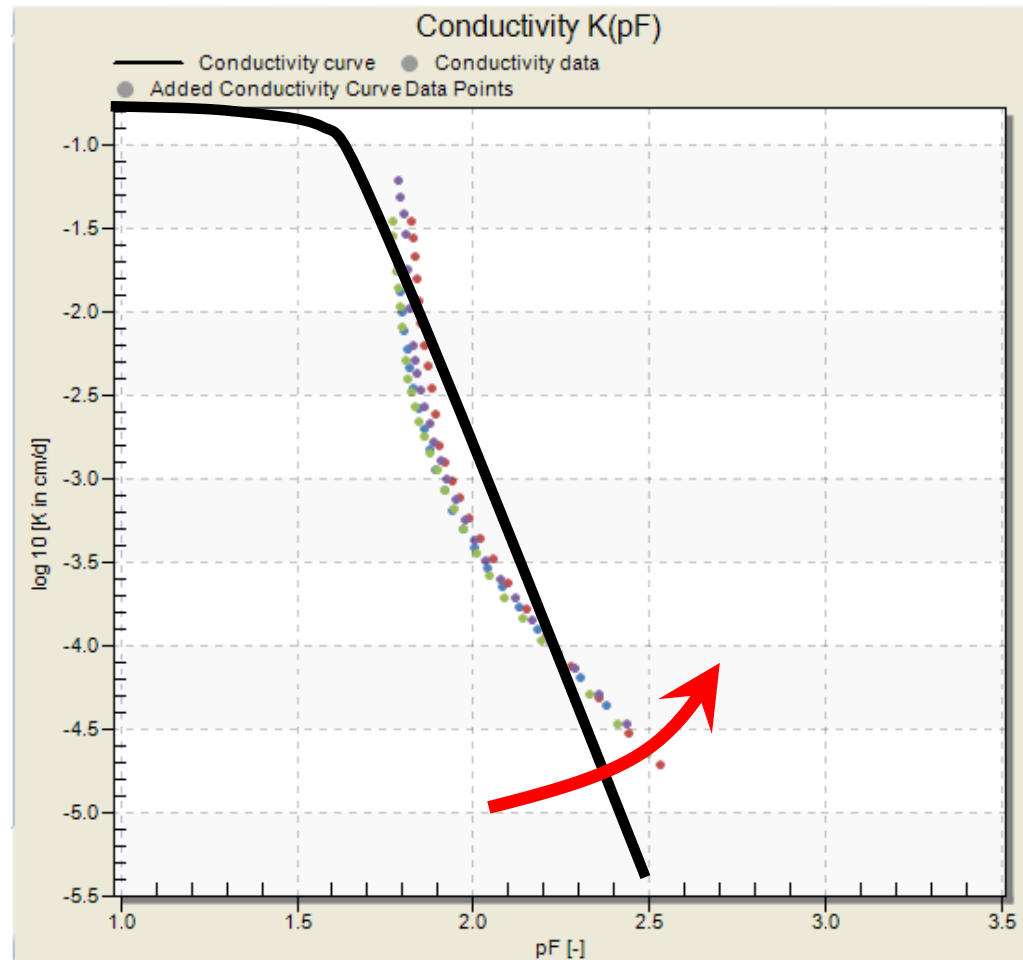
Theoretically: inverse of tortuosity!



$$0 < l < 1$$

De facto: empirical fitting parameter

Excursus 2: The tortuosity coefficient



Excursus 2: The tortuosity coefficient

Negative tortuosity coefficients are an **artefact** of an inadequate model structure



Conclusions



Conclusions (1): Where do we go?

- The limits of the usability(!) of Richards' equation are not yet explored.
- Many limitations can be overcome by adapting hydraulic functions.
- No replacement of the current standard, rather identifying situations where specific adaptations are needed.
- Limitations are not given by the modeling ability but by the inability to derive appropriate parameters.
- More and better experiments are needed.


Conclusion (2): SHP parametrizations

- Parameterizations of SHP should be as simple as possible – but not simpler
- Traditional SHP have structural deficits
- Conductivity function in the medium moist range is dominated by film/corner flow which can [under certain conditions] strongly affect the water flow
- Avoiding structural errors is of great importance in development of pedotransfer functions



Technische
Universität
Braunschweig

Institut für Geoökologie
Bodenkunde und Bodenphysik



Thank you obrigada

Parameterization of models for soil hydraulic properties: Challenges and recent advances

Wolfgang Durner

Sascha Iden, Andre Peters, Efstathios Diamantopoulos,
Kai Germer, Tobias Weber, Benedikt Scharnagl